#### NASA TECHNICAL NOTE



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DIFFUSION ALGORITHMS AND DATA REDUCTION ROUTINE FOR ONSITE LAUNCH PREDICTIONS FOR THE TRANSPORT OF TITAN III C EXHAUST EFFLUENTS

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prediction of the surface impact	for the dispersive transport of the	e exhaust efflue	nts from
the launch of a Titan III vehicle.	This specialization permits thes	e transport pre	dictions
to be made at the launch range i	n real time so that the effluent mo	nitoring teams	can
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# SYMBOLS AND DEFINITIONS

# Equations

D(x,y,z)	= dosage at the point x, y, and z (ppm-sec or mg sec/m <sup>3</sup> )
F	= buoyancy term in the instantaneous cloud rise formula $= \frac{3g Q_{H}}{4c_{p} \pi T_{s} \rho_{s}}$
	p s s
Н	= height of the stabilized exhaust cloud (meters)
K(r,t,p,T)	= diffusion coefficient
$\mathbf{L_{i}}$	= ith dimension of the rocket exhaust cloud (meters)
M	= molecular weight
$Q_{H}$	= heat release source strength (cal)
$Q_{\mathbf{M}}$	= mass source strength of the exhaust cloud (ppm)
R	= universal gas constant (0.289 Joules/gram deg K)
т	= temperature (degree Kelvin)
c p	= specific heat of air at constant pressure (0.24 cal/g deg K or 1.003 Joule/gram deg K)
${f f}$	= fractional amount of the total effluent which is released by the rocket in the surface mixing layer
g	= gravitational acceleration (9.8 m/sec <sup>2</sup> )
m	= power law exponent for the vertical profile of the wind azimuth.

= power law exponent for the wind speed.

- q = power law exponent for the vertical profile of the
- standard deviation of wind elevation angle in the surface mixing layer
- $r_{R}$  = initial cloud radius at the rocket exit
- s = stability parameter

р

$$= \frac{\mathbf{g}}{\mathbf{T}} \quad \frac{\partial \Phi}{\partial \mathbf{z}}$$

- t = time required for the exhaust cloud to reach equilibrium with the atmosphere at the stabilization height
- $\vec{u}$  = mean (time) wind speed (m/sec)
- <u>> = average (space) wind speed (m/sec)
- x = down range distance in the wind direction from the point of cloud stabilization (meters)
- y = distance from the centerline along the wind direction (meters)
- z = height of the stabilized exhaust cloud (meters)
- $\alpha$  = horizontal diffusion coefficient
- $\beta$  = vertical diffusion coefficient
- $\gamma$  = entrainment coefficient (Titan: 0.64)
- $\rho$  = density of the ambient air (mg/m<sup>3</sup>)
- $\sigma_{i}$  = standard deviation of the distribution of the exhaust effluents in the exhaust cloud in the ith direction (meters)

$$= \frac{L_{i}}{4.3}$$

$^{\sigma}$ AR	=	standard deviation of the wind azimuth angle at the surface
$^{\sigma}_{ m AT}$	=	standard deviation of the wind azimuth angle at the top of the layer
$\sigma_{ m ER}$	=	standard deviation of the wind elevation at the surface
$^{\sigma}_{ m  ET}$	=	standard deviation of the wind elevation angle at the top of the surface mixing layer
Φ	=	potential temperature
$\frac{\partial \Phi}{\partial \mathbf{z}}$	=	vertical gradient of the potential temperature $(\Phi)$
ΔΘ	=	change in wind direction between the top and bottom of the surface mixing layer
	=	$\Theta_{\mathbf{T}} - \Theta_{\mathbf{B}}$

## **Terms**

= the concentration (ppm or mg/m<sup>3</sup>)

Centerline:

 $\chi(\vec{r},t)$ 

The radial vector in the direction of the mean wind direction whose origin is the launch site.

Concentration:

is the amount of the effluent present at a specific time. The average concentration is the average amount present during the event.

Dosage: is the measure of the total amount of effluent (time

integrated concentration) due to the vehicle launch at a

specific location.

Ground Cloud: That cloud of rocket effluents emitted during the initial

phase of vehicle launch. This cloud is assumed to have

an ellipsoid shape.

Plume Cloud: The cloud of rocket effluents emitted from the vehicle in

flight. This cloud has a cylindrical shape whose height is

defined by the vertical thickness of the layer.

Potential Tem-

perature  $(\Phi)$ : is the temperature a volume of dry air would have if

brought adiabatically from its initial state to the standard

pressure of 100mb.

Quasiadiabatic

Layer: is a layer in which the vertical potential temperature

gradient is zero or less.

Stable Layer: is a layer in which the vertical potential temperature

gradient is positive.

#### Definition For Variable And Register Storage In The Program

#### 1. MET Routine

Lines  $0 \rightarrow 11$ 

Ro. R5. R10...R90 Altitude Data (meters or feet)

R1, R6, R11...R91 Wind Direction Data (Deg)

R2, R7, R12...R92 Wind Speed Data (m/sec or kn)

R3. R8. R13...R93 Temperature (Dry Bulb deg C)

R4, R9, R14...R94 Pressure (mb)

B increment for storage (by "5's") C increment for register location (transfer) X transfer increment

Y maximum transfer location

Z transfer limit

Lines  $12 \rightarrow 20$ 

Line 13 RX temperature data
RY, R(Y-5) altitude data
R(Y-1), R(Y+4) pressure data
R(Y+3), R(Y-2) temperature data

Line 18 RX wind speed data

Line 20 RX temperature data
R3 surface temperature dry bulb deg C
R28 temperature data
R97 Y position of the word 'TEMP-DRY''
R98 surface pressure mb
R99 surface temperature dry bulb deg C

B data conversion increment limit
X data increment
Y altitude calculation/conversion increment

Lines 21→26

R0 data count number
R1 surface temperature dry bulb deg C
R2 surface density GR/m<sup>3</sup>
R25 X position of the word "TEMP-DRY"

Line 21 A increment for register storage C transfer limit

Line 25 A transfer increment
C increment for register storage and limit

R4, R8, R12...R76 altitude data meters R5, R9, R13...R77 wind direction degrees R6. R10. R14...R78 wind velocity meters/sec R8, R11, R15...R79 potential temperature deg C Lines 27→58 R7 potential temperature at the surface deg C R16 Y position of the word 'wind speed' R18 X position of the word "wind speed" R24 Y position of the word "wind dir" R25 X position of the word "wind dir" R44 Y position of the word ''potential temp'' R47 X position of the word ''potential temp'' R88 maximum wind direction R89 minimum wind direction R90 month of meteorology data R91 day of meteorology data R92 year of meteorology data R93 time of meteorology data RX altitude data R(X+1) wind direction data R(X+2) wind speed data R(X+3) temperature data Line 37 RX wind direction data A X position of the word "adiabatic potential temp. grad." B Y position of the word "adiabatic potential temp. grad." Y adiabatic gradient increment Line 42 A conversion of maximum wind direction Line 43 B conversion of minimum wind direction Y differences of A and B C initial position for plotting wind direction scale

Line 46

X plotting increment

#### Line 48 X data increment

Line  $58 \rightarrow 75$ 

R1 Surface temperature deg K

R64 time data count

R76 R(X+4)-R(X-4)

R82 slope of temperature over altitude

R85 time of rise

R86 cloud height

R99 wind velocity at maximum cloud height

Line 71 R82 wind direction at maximum cloud height

Line 66 R85 time of maximum cloud rise

Line 66 R86 maximum cloud height

R(X+4), R(X-4), RX altitude data

R(X+5), R(X-3) wind direction data

R(X+6), R(X-2) wind speed data

R(X+7) temperature data

R(64+Z) time data

Line 68 R(64+Z) time of maximum cloud rise

A time square of adiabatic cloud rise model

B altitude increment

C potential temperature gradient

Y time of cloud rise

Z time storage increment

#### 2. Diffusion Routine

Lines  $0\rightarrow 12$ 

R1 surface temperature deg K

 $^{R1}\sigma_{X}$ 

```
Line 5 R2 surface pressure mb
```

R5 wind direction at the surface

R77  $\sigma_{_{
m f V}}$ 

 $R78Q_{K}$ 

R79 wind direction at the top of the layer

R80 maximum concentration

R81 wind speed at the top of the layer

R83  $\sigma_{
m Ar}$ 

 $R84 \sigma_{ATK}$ 

R85 time of maximum cloud rise

R86 maximum cloud H

R87 top of the layer

R88 concentration

R89 radius of cloud

R90 month of meteorology data

R91 day of meteorology data

R92 year of meteorology data

R93 time of meteorology data

R97 dummy

R98 down range distance of maximun concentration

 $R101 \sigma'_{AK} = \sigma'_{EK}$ 

R102 plotter constant

R103 plotter constant

 $A (\sigma_{Z}^{2})/2$ 

B verticle component of multilayer diffusion model

C power law exp. - verticle profile of deviation

X down range distance

Y dosage

#### Line 6 A concentration

Lines  $13 \rightarrow 17$ 

R76 R(X+4)-R(X-4)

R(X+4), R(X-4), RX altitude data

R(X+5), R(X-3), R(X+1) wind direction speed R(X+6), R(X-2), R(X+2) wind speed data X data increment

Lines 18→30

R64 initial time location

R76 X position at maximum cloud rise

R77 Y position at maximum cloud rise

R82 wind direction at maximum cloud rise

R94 X position at maximum cloud rise

R95 Y position at maximum cloud rise

R96 time after maximum cloud rise at an arbitrary distance

R97 time count limit

R99 wind speed at maximum cloud rise

Line 26 R76 X position at an arbitrary down range distance after maximum cloud rise

R77 Y position at an arbitrary down range distance after maximum cloud rise

A polar position of X value of plotter

B time increment

C down range distance

Y polar position of Y value of plotter

Z down range angle

Lines  $31 \rightarrow 49$ 

R8 plotter constant

R9 plotter constant

R17 plotter constant

R18 plotter constant

R20 Y test value for zero in Y

R21 X test value for zero in Y

R22 X test value for zero in Y

R23 Y test value for zero in Y

R50 Y coordinate in plotter R51 -Y coordinate in plotter R52 X coordinate in plotter R53 -X coordinate in plotter X down range distance Y cross range distance

#### FA subroutine

P1 down range distance

P2 cross range distance

P3  $\sigma_{Y}$ 

 $^{P4}\sigma_{X}$ 

P5 vertical component of multilayer diffusion model

R0 L(XK)

R1  $\sigma_X$ 

R2 uK

R5 wind direction at the surface

R6 wind speed at the surface

R77  $\sigma_{Y}$ 

 $R78Q_{K}$ 

R79 wind direction at the top of the layer

R81 wind speed at the top of the layer

R86 height of maximum cloud rise

R87 height of layer

R88 concentration

R89 radius of the cloud

R101  $\sigma'_{EK} = \sigma'_{AK}$ 

A  $(\sigma_Z)^2/2$ 

B summation on multilayer diffusion model Z i

Line 4 B vertical component of multilayer diffusion model

Line 6 Z power law exp. for wind speed profile in surface layer

# FOR ONSITE LAUNCH PREDICTIONS FOR THE TRANSPORT OF TITAN III C EXHAUST EFFLUENTS

#### **SUMMARY**

Specialization of the NASA/MSFC Multilayer Diffusion Model for the Titan III C vehicle permits us to utilize a programmable calculator (HP9820A) at the launch site in making a forecast of the surface effects from the exhaust effluents of this vehicle.

This specialization of the general diffusion model limits the scope of analysis to the surface mixing layer with the ellipsoidal dispersion of the exhaust cloud (Model 3) — the resulting concentration and dosage prediction are limited to the earth's surface.

The only mandatory inputs to the Titan III C version of the NASA/MSFC Surface-Layer Diffusion Model are the data points from a rawinsonde sounding and the altitude of the top of the surface mixing layer, because of the specialization. The data reduction routine employs the calculator plotter to display the meteorological profile, the temporal history of the rise of the Titan exhaust cloud to the point of stabilization, and a ground level delineation of the concentrations and dosage of HCl along the exhaust cloud path. In addition, the HCl concentration isopleths are plotted on a map of Kennedy Space Center and Cape Canaveral.

The NASA/MSFC Surface Layer Diffusion Calculator Routine for the Titan III C is designed for both real-time, launch-site predictions and as a guide to determining the inputs for the NASA/MSFC Multilayer Diffusion Computer Program.

#### SECTION 1. INTRODUCTION

The NASA/MSFC Multilayer Diffusion Algorithms have been specialized for the Titan III C vehicle into the NASA/MSFC Surface Layer Diffusion Calculator Program to provide launch-site predictions of the rocket exhaust effluent transport. This calculator program has been designed for launches at Kennedy Space Center.

The National Environmental Policy Act of 1969 and the April 23, 1972, guidelines of the Council on Environmental Quality require impact statements for assessing the environmental impact of the Space Shuttle and other NASA space vehicle rocket motor effluents. Development of quantitative procedures for estimating the space vehicle rocket motor exhaust effluents hazard has been underway for over a decade at the NASA's Marshall Space Flight Center. These computerized procedures for estimating the tropospheric transport of potentially toxic exhaust effluents have been developed utilizing a Eulerian model [1]. In addition to estimates of atmospheric transport, dispersal, and decay of all airborne toxic material released as a result of normal launch operations, estimates must also be provided for cases involving fuel spillage, vehicle abort, or vehicle destruct situations.

Universally accepted and adequately validated prediction techniques for the rocket motor effluent problem are not available, and some uncertainty exists concerning very important aspects of this problem, such as the amount and composition of the rocket motor effluents and their dispersal and transport in the atmosphere. The available atmospheric measurements to ascertain the reliability of the description of rocket effluent dispersion models in the atmosphere are sparse and of questionable accuracy. On the other hand, the requirements for estimating toxic fuel hazards clearly exist in order to establish special constraints on operations, test, and launch activities to assure allowable concentrations of these effluents will not be exceeded. The need for implementing a program for monitoring rocket engine exhaust effluents has been recognized for many years. As a result of informal discussions between representatives of NASA Headquarters, Marshall Space Flight Center, Langley Research Center, and Kennedy Space Center, it became apparent that a NASA inhouse rocket motor effluent prediction and measurement program was desirable, possible, and feasible.

A joint solid rocket motor exhaust prediction (Marshall) and measurement (Langley supported by Kennedy) program has evolved utilizing the Titan launches as a source for empirical information that can be employed to

more accurately predict the environmental impact of the Space Shuttle under varying atmospheric conditions. In order to determine the locations and the operating duration of the sensors, in the measurement part of the program, the NASA/MSFC Multilayer Diffusion Model is used to predict the ground-level concentrations and dosage; otherwise, a vast network of expensive sensors would be needed. In addition to reducing the cost of the measurement program, the transport predictions are necessary to determine the proper empirical constants for the model. Thus real-time onsite effluent transport predictions are desirable, if feasible. This feasibility implied that a mobile substitute for the computer is needed to transform the description of the atmospheric kinematics and thermodynamics obtained from rawinsonde soundings and meteorological towers into a predicted effluent transport description.

The substitute for the computer utilized is a programmable calculator (HP9820A). Naturally, this means that the computer program, which requires a core storage of about 30K, must be reduced in size. This was achieved by taking the NASA/MSFC Multilayer Diffusion Model — which provides a transport description for any rocket exhaust effluents from any vehicle, anywhere in the troposphere, under any environmental conditions, and for different types dispersive — and limiting the description's applicability to a specific effluent — HC1, to a specific vehicle — Titan III C, to a specific layer in the troposphere — surface mixing layer, and to spherical diffusion without deposition effects. The requirement for real-time predictions led to the incorporation of meteorological and buoyancy algorithms as an integral part of the NASA/MSFC Surface Layer Diffusion Program. In addition, user requirements dictated that the surface concentration footprint of the exhaust cloud be superimposed on a map of the Kennedy Space Center area.

The modeling considerations utilized to achieve these results are discussed in the next section. The computational procedures and graphical results are discussed and provided in the third section.

#### SECTION II. NASA/MSFC SURFACE-LAYER DIFFUSION MODEL

The spatial description, in terms of concentration and dosage, of the dispersive transport of effluents from a discrete source is afforded by the NASA/MSFC Multilayer Diffusion Model. Specifically, this application of the model is for the prediction of the distribution of the toxic effluents [hydrogen chloride (HCl), carbon monoxide (CO), carbon dioxide (CO<sub>2</sub>), and alundum (Al<sub>2</sub>O<sub>3</sub>)] associated with the rocket exhaust emitted during the launch of a

Titan III vehicle from Cape Canaveral, in order to assess the resulting environmental effects. The dispersive description accorded by this Multilayer Diffusion Model is initiated at the point where the rocket exhaust cloud of effluents reaches thermodynamic equilibrium with the environment—cloud stabilization—and therefore depends strongly on the kinematic and thermodynamic profiles of the atmosphere, along with the chemical and thermodynamic composition of the exhaust cloud.

Thus, the initial considerations in this section are given to the atmospheric description employed as the input to the cloud rise algorithms. The discussion of the stabilized cloud considers the algorithms for both the cloud rise and cloud dimensions utilized as the initial inputs to NASA/MSFC Multilayer Diffusion Model. The final part of this section provides the significant mathematical expressions that were employed to obtain the center line concentrations and dosages, and the concentrations isopleths.

## A. Meteorological Profile

The influence of the atmospheric conditions on the dispersive transport of rocket exhaust effluents is the basis for the format of meteorological profile.

To illustrate, consider the diffusion equation. The solution of the differential equation for the diffusion dictates that we treat the diffusion problem in two stages [2]. The nonlinear differential form of the diffusion equation is

$$\frac{\partial \chi(\overrightarrow{\mathbf{r}},t)}{\partial t} + \langle \overrightarrow{\mathbf{v}} \rangle \cdot \nabla \chi(\overrightarrow{\mathbf{r}},t) = \nabla \cdot [\widetilde{K}(\overrightarrow{\mathbf{r}},t,p,T) \cdot \nabla \chi(\overrightarrow{\mathbf{r}},t)] ,$$
(1)

where:

 $\chi(\vec{r},t)$  is the scalar concentration of the diffusing gas,

<v> is the mean wind velocity, and

 $\widetilde{K}(\overrightarrow{r},t,p,T)$  is the diagonal diffusion tensor which is a function of position, time, and the thermodynamic parameters, pressure and temperature.

To enable us to solve this differential equation by the separation of variables, we must establish the necessary assumptions to make this equation linear. The normal technique—and the one used in the NASA/MSFC Multilayer Diffusion Model—to linearize the diffusion equation (Equation 1) is by restricting our model to a kinematic description by assuming that the diffusion coefficient (K) is a time average value and thermodynamically independent. This implies the initial conditions occur when the rocket exhaust cloud achieves thermodynamic equilibrium with the atmosphere at cloud stabilization. Consequently, the effluent transport problem can be decomposed into two phases: the thermodynamic phase during cloud rise and the kinematic phase of diffusion.

The atmospheric thermodynamic parameters (pressure, temperature, and density) govern the magnitude of the buoyant force on the exhaust cloud and thus dictate the height of cloud rise. Incorporating these atmospheric thermodynamic parameters into a suitable concise thermal description that will efficiently interface with the cloud rise algorithms requires the following considerations. Since the temperature (T) is a function of the pressure (p), as expressed in Poisson's equation

$$\frac{T}{T_o} = \left(\frac{p}{p_o}\right)^{\frac{R}{C}} p , \qquad (2)$$

where:

R is the universal gas constant,

 ${\displaystyle \mathop{\mathbf{C}}_{\mathbf{p}}}$  is the specific heat at a constant pressure, and

$$R/C_p = 0.288 .$$

The concept of a potential temperature ( $\Phi$ ) is introduced to reference the temperature to a specific pressure (1000 mb) and is defined as

$$\Phi = T(\frac{1000}{p})^{0.288} . (3)$$

The potential temperature can be shown to be a measure of the entropy ( $s = C \ln \Phi + Const$ ), therefore, the vertical potential gradient ( $\partial \Phi / \partial z$ ) is a measure of the change in entropy. Since in an adabatic process there is not a change in entropy, the potential temperature gradient is zero (This corresponds to a straight, vertical line on our meteorological profile).

In order to achieve exhaust cloud stabilization with the atmosphere, we must achieve an entropy balance between the exhaust cloud and the atmosphere, which can be determined by utilizing the thermodynamic description afforded by the potential temperature profile. In the case of a hot rocket exhaust cloud, this balance results from both entrainment due to the turbulence mixing of this cloud and the exhaust cloud rising in the atmosphere to a region of higher entropy. If the potential temperature difference between the surface and a cloud height is negative or zero — what we shall define here as an adiabatic condition, the entropy difference between the exhaust cloud and atmosphere will continue to increase and cloud stabilization will not occur. However, if the potential temperature gradient is positive - a stable condition - the entropy difference between the exhaust cloud and the atmosphere decrease as the exhaust cloud rises until equilibrium is obtained. Thus, the thermodynamic influences of the atmosphere on the hot rocket exhaust cloud during the initial transport stage where the exhaust cloud is rising to the point of equilibrium can be determined directly from the potential temperature profile.

During the kinematic stage of the rocket exhaust effluent transport, the potential temperate profile can also be a guide to the top of the surface mixing layer in that the signature for the top of the surface mixing layer is characterized by a change in the meteorological conditions. Specifically, this signature can be a change of wind velocity or a change in thermal gradient or both. The change in thermal gradient is normally due to a temperature inversion or possibly an isothermal layer above the altitude of cloud stabilization — both of which result in a more positive potential temperature gradient.

The NASA/MSFC Multilayer Diffusion Model depends strongly on the wind direction and speed profiles in the surface mixing layer (first 2 kilometers) to determine the surface effects from the kinematics of the exhaust cloud. These profiles are required not only for the mean wind speed and direction at the cloud stabilization height, but these profiles are also required for the wind speed and directional shears over the surface mixing layer.

The meteorological profile (wind speed, wind direction, temperature, and pressure) is normally obtained from a rawinsonde sounding of the atmosphere. To obtain the entropy profile required for these soundings, we translate the dry bulb temperature and pressure into the potential temperature in accord with equation (3).

If the levels of the sounding are referenced only in terms of pressure, rather than altitude (z) and pressure, we compute the altitude of the level using

$$z = 29.3 (T + T_0) ln \frac{p_0}{p}$$
 , (4)

where the subscript "o" refers to the next lower level which is referenced in the sounding to an altitude.

Thus the atmospheric thermodynamic parameters required for the cloud rise plume are the pressure and temperature which are utilized in obtaining the potential temperature profile. The atmospheric kinematic parameters required for the diffusion phase are the wind speed and direction. The delineation of these parameters as a function of altitude comprise the meteorological profile.

#### B. Titan III C Exhaust Cloud-Rise Relation

During a normal Titan III C launch, the burning of the rocket motors results in the formation of a hot exhaust cloud of the rocket exhaust effluents. Subsequently, this exhaust cloud rises and entrains ambient air until thermodynamic equilibrium between the exhaust cloud and the ambient atmosphere is attained. The Titan employs solid rocket motors which are characterized by a rapid maximum thrust after ignition; therefore, the Titan has a relatively short residence time on the pad after ignition. This characteristic enables us to assume spherical entrainment of the ambient air; that is, the ground exhaust cloud of the Titan is treated as an instantaneous source. Based on these assumptions, we will develop the cloud rise equations for the Titan III C rocket which will afford similar results to those obtained by G. A. Briggs [3]. The description developed here for the temporal delineation of the cloud rise during the thermodynamic stage is designed to not only afford the maximum height, but this description also affords the temporal delineation of cloud rise. This temporal delineation of the cloud rise is in turn coupled with the atmospheric kinematic profile to obtain the spatial history of the cloud rise in order to locate the point of the exhaust cloud stabilization.

The cloud rise formulation, as was pointed out in the last section, is dependent on the stability of the thermodynamic condition in the atmosphere. The potential temperature gradient  $(d\Phi/dz)$  is the index of the stability, which is given by

$$\frac{d\Phi}{dz} = \left(\frac{1000}{p}\right)^{\frac{R}{c}} p \frac{\partial T}{\partial z} + \frac{g}{c} = \frac{\Phi}{T} \frac{\partial T}{\partial z} + \frac{g}{c} , \qquad (5)$$

where the potential temperature  $(\Phi)$  is defined in Equation (3). The other variables in Equation (4) are the dry bulb temperature (T:deg K), the pressure (p: millibars), and the altitude above the surface (z: meters). The constants are the specific heat of air at constant pressure ( $C_p = 1.003$  Joules/gram deg K), the gas constant (R = 0.289 Joules/gram deg K), and the gravitational acceleration (g: m/sec²). If

$$\frac{\Delta \Phi}{\Delta z} \leq 0 \quad , \tag{6}$$

where  $\Delta\Phi$  and  $\Delta z$  are the potential temperature difference and the altitude difference between the surface and the altitude of interest, then we will assume adiabatic rather than stable conditions exist. The cloud rise height (z) for a Titan III for an adiabatic atmosphere is given by [4],

$$z_{A} = \left[ \frac{2 \text{ Ft}^{2}}{\gamma^{3}} + \left( \frac{r_{R}}{\gamma} \right)^{4} \right]^{\frac{1}{4}} - \frac{r_{R}}{\gamma} ; \qquad (7)$$

whereas, the cloud rise height  $(z_s)$  for a Titan III for a stable atmosphere is given by [4]

$$z_{s} = \left\{ \frac{4 w_{o} r_{R}^{3}}{\gamma^{3} s^{\frac{1}{2}}} \sin \left( s^{\frac{1}{2}} t \right) + \frac{4F}{\gamma^{3} s} \left[ 1 - \cos \left( s^{\frac{1}{2}} t \right) \right] + \left( \frac{r_{R}}{\gamma} \right)^{4} \right\}^{\frac{1}{4}} .$$
(8)

The variables are the initial vertical cloud speed ( $w_c$ : m/sec), the buoyance term (F:m<sup>4</sup>/sec<sup>2</sup>), the stability term (s: sec<sup>-2</sup>), and the time after ignition (t: sec). The constants are the entrainment coefficient ( $\gamma = 0.64$ ) and the initial cloud radius ( $r_R$ : m). For the Titan, we will assume the initial cloud radius ( $r_R$ ) is zero, which permits us to rewrite Equations (7) and (8) as:

(1) Adiabatic Case -

$$t_A = \sqrt{\frac{z^4 \gamma^3}{2F}} \equiv \sqrt{\alpha}$$
 , and (9)

(2) Stable Case -

$$t_{s} = \frac{\arccos\left(1 - \frac{s\alpha}{2}\right)}{\sqrt{s}} . \tag{10}$$

The buoyance term (F) is

$$F = \frac{3g Q_H}{4 c_p \pi \rho_s T_s} , \qquad (11)$$

where:

 $\rho_{_{\rm S}}$  (gm/m³) is the surface density,

 $\boldsymbol{T}_{_{\boldsymbol{S}}}$  (°K) is the surface temperature, and

 $Q_{H}$  (cal) is the effective heat release.

Based on a least square curve fit of Titan III C data [4], the heat release as a function of cloud height is

$$Q_{H} = 1.609907 \times 10^{9} z^{0.4837}$$
, (12)

where we assume that the flame trench is filled with water.

The temporal cloud rise paths given in Equations (9) and (10) can now be written as:

#### (1) Adiabatic condition for Titan III C -

$$t_{A} = 2.889827772 \times 10^{-6} z^{1.75815} \sqrt{T_{s} \rho_{s}}$$
 (13)

#### (2) Stable condition for Titan III C -

$$t_s = \arccos \left[1 - 4.175552276 \times 10^{-12} \text{ s } T_s \rho_s z^{3.5163}\right],$$
 (14)

where the stability term is

$$s = \frac{g}{T_s} \frac{\Delta \Phi}{\Delta z} = \frac{9.8}{T_s} \frac{\Delta \Phi}{\Delta z} . \qquad (15)$$

The potential temperature gradient is taken directly from the potential temperature profile at launch time.

The exhaust cloud rise algorithm for the Titan III C (Equation 13 and 14) affords only a temporal description of the ascent of the exhaust cloud to the point of cloud stabilization. Examination of the adiabatic cloud rise relation (Equation 13) reveals that a time for cloud stabilization is not defined, and this may be disturbing to some. However, two points should be kept in mind:

(1) When there is an adiabatic lapse rate in the atmosphere, this lapse rate exists in a relatively thin layer near the ground, with a stable layer above it.

(2) Since, as was pointed out in the last part, an entropy balance between the exhaust cloud and its local environment is required, cloud stabilization only can occur where there is a positive potential temperature difference between an altitude and the surface, that is, in a stable layer. The altitude range of the stable cloud rise algorithm (Equation 14) is between the surface and cloud stabilization; i.e.,

$$0 \le z \le 2.097795927 \times 10^3 \left(\frac{T_s \rho_s}{s}\right)^{0.28439}$$
 (16)

Beyond these limits, the periodic nature of the algorithm results in variations outside of these limits. Since the stability term (s) is a function of altitude (Equation 15), the maximum cloud rise height requires some form of an iterative solution to obtain this height.

To summarize, the temporal cloud rise delineation that is obtained from the cloud rise algorithms is interfaced with the meteorological kinematic to predict the cloud path before stabilization. It should be noted that the cloud rise algorithms given in this part are valid only for the Titan III since the source strength utilized is for the Titan.

# C. Titan III Version of NASA/MSFC Multilayer Diffusion Model

By specialization of the NASA/MSFC Multilayer Diffusion Model for Titan III exhaust effluents, the prediction for the ground level concentration isopleths can be obtained from a small programmable desk calculator (HP9820), in real time. The modeling approach employed is as follows.

The general differential equation for kinematic diffusion (Equation 1) can be linearized by assuming that the meteorological profile represents the average atmospheric conditions over the area of interest and solved by separation of variables for the spatial distribution of the concentration and dosage resulting from the launch of Titan III. The temporal variations are accounted for by the standard deviation terms for the elevation and azimuth. Because of the complexity of the resulting formulation, we will first present a generalized model as an introduction to the actual algorithm [1].

The generalized concentration model for a nearly instantaneous source is expressed as the product of five modular terms:

whereas, the generalized dosage model for a nearly instantaneous source is defined by the product of four modular terms:

Dosage = { Peak Dosage Terms} 
$$\times$$
 { Lateral Term}  $\times$  { Vertical Term}  $\times$  { Depletion Term} .

Thus, the mathematical description for the concentration and dosage models permit flexibility in application to various sources and for changing atmospheric parameters while always maintaining a rigorous mass balance.

Two obvious differences exist. First, the peak concentration term refers to the concentration at the point x, y = 0, z = H (where x is the wind direction and H is any height) and is defined by the expression

Peak Concentration = 
$$\frac{Q_{M}}{(2\pi)^{3/2}\sigma_{x}\sigma_{y}\sigma_{z}},$$
 (17)

where  $\mathbf{Q}_{\mathbf{M}}$  is the mass source strength and  $\sigma_{\mathbf{i}}$  is the standard deviation of the concentration distribution in the  $\mathbf{i}^{\mathbf{th}}$  direction. The peak dosage term is given by

Peak Dosage = 
$$\frac{Q_{M}}{2\pi \overline{u} \sigma_{y}^{\sigma} z}$$
, (18)

where u is the mean wind speed over the layer. The second difference between these models is that the concentration contains a modular alongwind term to account for downstream temporal effects not considered in the dosage model. The alongwind term affords an exponential decay in concentration as a function of: cloud transit time, concentration distribution and the mean wind speed.

The <u>lateral</u> term (which is common to both models) is another exponential decay term and is a function of the Gaussian spreading rate and the distance laterally from the mean wind azimuth. The <u>vertical term</u> (again common to both models) is a rather complex decay function since it contains a multiple reflection term for the point source which stops the vertical cloud development at the top of the mixing layer and eventually changes the form of the vertical concentration distribution from Gaussian to rectangular. The last modular in both models is the <u>depletion term</u>. This term accounts for the loss of material by simple decay processes, precipitation scavenging, or gravitational settling. The depletion term will be neglected in the Titan III version of the model.

The meteorological profile is utilized in layering the atmosphere in accord with homogeneous kinematic and thermodynamic properties—hence the name ''Multilayer Diffusion Model''. The specialization of the general NASA/MSFC Multilayer Diffusion Model has limited the number of layers for consideration to just the surface mixing layer. In addition, we assume that the source has a spherical shape with ellipsoidal expansion (Model 3 of the NASA/MSFC Multilayer Diffusion Model [1]).

Thus the dosage algorithm is

$$D\left\{x,\,y,\,z_{B}^{}\right. < z^{}\right. < z_{T}^{}\right\} = \frac{Q_{M}^{}}{2\pi\,\sigma_{y}^{}\,\sigma_{z}^{}\,\bar{u}} \qquad \left\{\exp\left[\frac{-y^{2}^{}}{2\sigma_{y}^{2}^{}}\right]\right\} \left(\exp\left[\frac{-\left(H^{}-z\right)^{2}^{}}{2\sigma_{z}^{2}}\right]\right\}$$

$$+ \exp \left[ \frac{-\left( {\rm H} - 2z_{\rm B} + z \right)^2}{2\sigma_{\rm z}^{\ 2}} \right] + \sum_{i=1}^{\infty} \left\{ \exp \left[ \frac{-\left[ 2i \left( z_{\rm T} - z_{\rm B} \right) - \left( {\rm H} - 2z_{\rm B} + z \right) \right]^2}{2\sigma_{\rm z}^{\ 2}} \right] \right.$$

$$+ \exp \left[ -\left[ \frac{2i \left(z_{T} - z_{B}\right) + \left(H - z\right) \right]^{2}}{2\sigma_{z}^{2}} \right] + \exp \left[ -\left[ \frac{2i \left(z_{T} - z_{B}\right) - \left(H - z\right) \right]^{2}}{2\sigma_{z}^{2}} \right]$$

$$+ \exp \left[ -\left[ \frac{2i \left(z_{T} - z_{B}\right) + \left(H - 2z_{B} + z\right) \right]^{2}}{2\sigma_{z}^{2}} \right] \right\} \right) , \qquad (19)$$

where:

 $\boldsymbol{Q}_{\boldsymbol{M}}$  corresponds to the source strength or total mass of material in the layer,

H is the height of the centroid of the stabilized cloud, and subscripts T and B stand for the top and bottom.

By restricting the dosage mapping to the surface and defining the bottom of the layer as the surface, Equation (18) simplifies to

$$D\{x, y, o\} = \frac{Q}{2 \pi \sigma_{y} \sigma_{z} \overline{u}} \left[ \exp \left\{ \frac{-y^{2}}{2 \sigma_{y}^{2}} \right\} \right] \left[ \exp \left\{ \frac{-H}{\sigma_{z}^{2}} \right\} \right]$$

$$+ \sum_{i=1}^{n} \left[ \exp \left\{ -\frac{\left(2i z_{T} + H\right)^{2}}{\sigma_{z}^{2}} \right\} + \exp \left\{ -\frac{\left(2i z_{T} - H\right)}{\sigma_{z}^{2}} \right\} \right]$$

$$(20)$$

where  $\mathbf{z}_t$  is the altitude of the top of the surface mixing layer and n is such that the first exponential in the summation is greater than 100. This is the specialized dosage algorithm that we use.

The source strength (Q:ppm) for the HCl from the Titan III is

$$Q = 4.37858 \times 10^8 \frac{T_s}{P_s} \times H^{0.4,837} , \qquad (21)$$

where the surface temperature (T; degree K) and the surface pressure (p;:mb) are used, since we are interested only in the concentrations and dosages at the surface.

The standard deviation of the crosswind distribution  $(\sigma_y)$ , the standard deviation of the vertical distribution  $(\sigma_z)$ , and the mean wind speed  $(\bar{u})$  are defined as:

a. The standard deviation of the crosswind dosage distribution is defined by

$$\sigma_{y} = \left\{ \left[ \sigma_{A}^{\dagger} \left\{ \tau \right\} \right] \times_{ry} \left( \frac{x + x - x (1 - \alpha)}{\alpha x_{ry}} \right)^{\alpha} \right]^{2} + \left[ \frac{\Delta \theta^{\dagger} x}{4 \cdot 3} \right]^{2} \right\}$$

$$+ \left[ \frac{\Delta \theta^{\dagger} x}{4 \cdot 3} \right]^{2}$$

$$(22)$$

where  $\sigma_A^{\,\prime}(\tau)$  corresponds to the mean layer standard deviation of the wind azimuth for the cloud stabilization time  $(\tau)$ . The difference in wind direction  $(\Delta\,\theta^{\,\prime}: {\rm radians})$  is taken between the surface and the top of the surface mixing layer in accord with

$$\Delta\theta' = (\theta_{\rm T} - \theta_{\rm B}) \left(\frac{\pi}{180}\right), \tag{23}$$

where  $\theta_T$  and  $\theta_B$  are the mean wind direction in degrees at the top and at the base of the layer, respectively. This is the wind shear. If we again assume that diffusion coefficient is one  $(\alpha = 1)$ , then Equation (22) becomes

$$\sigma_{y} = \left\{ \left[ \sigma_{A}^{\dagger} \left\{ \tau \right\} \left( x + x_{y} \right) \right]^{2} + \left[ \frac{\Delta \theta^{\dagger} x}{4.3} \right]^{2} \right\}^{\frac{1}{2}} . \tag{24}$$

From the relation, we can observe how important a factor the wind shear is in determining the crosswind distribution of the effluent. In the surface layer

$$\sigma_{A}^{\prime} \{ \tau \} = \frac{\sigma_{AR}^{\prime} \{ \tau \} \left[ (z_{T})^{m+1} - (z)^{m+1} \right]}{(m+1) (z_{T} - z) (z)^{m}}, \qquad (25)$$

where the standard deviation of the wind azimuth angle [ $\sigma_{AR}^{\dagger}(\tau)$ ] at height  $z_{R}$  and for the cloud stabilization time  $\tau$  is

$$\sigma_{AR}^{\dagger} \left\{ \tau \right\} = \sigma_{AR}^{\dagger} \left\{ \tau_{o} \right\} \left( \frac{\tau}{\tau_{o}} \right)^{-1/5} \left( \frac{\pi}{180} \right) . \tag{26}$$

Here  $\sigma_{AR}^{\ \ \{\tau_o\}}$  is the standard deviation of the wind azimuth angle at height  $z_R$  and for the reference time period  $(\tau_o)$ , and the power-law exponent (m) for the vertical profile of the standard deviation of the wind azimuth angle in the surface layer is

$$m = \log \left( \frac{\sigma_{AT}^{\dagger} \{ \tau \}}{\sigma_{AR}^{\dagger} \{ \tau \}} \right) \log \left( \frac{z_{T}}{z_{R}} \right) . \tag{27}$$

Then

$$\sigma_{\text{AT}}^{\prime} \left\{ \tau \right\} = \sigma_{\text{AT}}^{\prime} \left( \tau_{\text{O}}^{\prime} \right) \left( \frac{\tau}{\tau_{\text{O}}^{\prime}} \right) \left( \frac{\pi}{180} \right) , \qquad (28)$$

where  $\sigma_{AT}^{\{\tau_o\}}$  is the standard deviation of the wind azimuth angle at the top of the surface layer for the reference time period.

The standard deviation of the wind azimuth angle cannot be determined at the top of the surface mixing layer because a temporal history of the wind azimuth cannot be obtained. We normally calculate this using the power law relation

$$\sigma_{AT} \{ \tau_{o} \} = \sigma_{AR} \{ \tau \} \left( \frac{z_{T}}{z_{R}} \right)^{-P}$$
 (29)

where

$$P = \frac{\log\left(\frac{u_T}{u_R}\right)}{\log\left(\frac{z_T}{z_R}\right)},$$
(30)

where u is the wind speed at the top of the surface mixing layer.

The crosswind virtual distance is

$$x_{y} = \frac{\sigma_{y0}}{\sigma_{A}^{\prime} \{\tau\}} - x_{Ry} , \qquad (31)$$

where  $\sigma_{yo}$  is the standard deviation of the lateral source dimension, which is

$$\sigma_{\text{yo}} = \gamma H = 0.64 H , \qquad (32).$$

where H(m) is the height of the exhaust cloud stabilization.

b. The standard deviation of the vertical dosage distribution is defined by the expression

$$\sigma_{z} = \sigma_{E}' \times_{rz} \left[ \frac{x + x_{z} - x_{rz} (1-\beta)}{\beta x_{rz}} \right]^{\beta}, \qquad (33)$$

where:

 $\sigma_{E}^{t}$  describes the mean standard deviation of the wind elevation angle,

 $\mathbf{x}_{\mathbf{z}}$  gives the vertical virtual distance,

 $\beta$  accounts for vertical diffusion, and

x is the distance over which rectilinear vertical expansion occurs downwind from an ideal point source.

In this specialization, we assume that the vertical diffusion coefficient is one  $(\beta=1)$  which permits us to rewrite equation (33) as

$$\sigma_{\mathbf{z}} = \sigma_{\mathbf{E}}^{\dagger} \left( \mathbf{x} + \mathbf{x}_{\mathbf{z}} \right) \tag{34}$$

where

$$\sigma_{\mathbf{E}}^{\dagger} = \sigma_{\mathbf{A}}^{\dagger} \quad . \tag{35}$$

Thus,  $\sigma_A^{\,\prime}$  is obtained from Equation (30), since (in general) the value for  $\sigma_F^{\,\prime}$  cannot be obtained without special instrumentation.

The vertical distance  $\mathbf{x}_{\mathbf{z}}$  is given by the expression

$$x_{z} = \frac{\sigma_{zo}}{\sigma_{E}^{\prime}} - x_{Rz}$$
 (36)

where  $\sigma_{zo} = \sigma_{yo}$  is the standard deviation of the vertical dosage distribution at  $x_{Rz}$ , the distance from the source where the measurement is made in the surface mixing layer.

c. The mean speed of cloud transport  $(\bar{u})$  in the surface layer is defined in accord with the power-law

$$\bar{u}\{z\} = \bar{u}_R \left(\frac{z}{z_R}\right)^p$$
, (37)

where  $\bar{u}_R$  is the mean wind speed measured at the reference height  $z_R$  and the power-law exponent (p) for the wind speed profile in the surface layer is described by

$$p = \log\left(\frac{\bar{u}_{T}}{\bar{u}_{R}}\right) / \log\left(\frac{z_{T}}{z_{R}}\right) . \tag{38}$$

Here,  $\bar{u}_T$  corresponds to the mean wind speed at the top of the surface layer  $(z_T)$ . Thus, in the surface layer, the mean cloud transport speed  $(\bar{u})$  is

$$\bar{\mathbf{u}} = \frac{\bar{\mathbf{u}}_{\mathbf{R}}}{(\mathbf{z}_{\mathbf{T}} - \mathbf{z}_{\mathbf{R}}) \mathbf{z}_{\mathbf{R}}^{\mathbf{p}}} \int_{\mathbf{z}_{\mathbf{R}}}^{\mathbf{z}_{\mathbf{T}}} \mathbf{z}^{\mathbf{p}} d\mathbf{z} , \qquad (39)$$

which reduces to

$$\bar{u} = \frac{\bar{u}_{R}}{(z_{T} - z_{R})(z_{R})^{p}(1+p)} . \tag{40}$$

The concentration  $(\chi)$  follows directly from the results for dosage (D) algorithm given Equation (19). The average concentration then is just

$$\bar{\chi} \{x, y, z\} = D\{x, y, z\} \left[\frac{\bar{u}}{4.3 \sigma_{x}}\right],$$
 (41)

where the standard deviation of the alongwind concentration distribution (  $\sigma_{_{\rm X}}\!)$  in the layer is

$$\sigma_{X} = \left\{ \left[ \frac{L(X)}{4.3} \right]^{2} + \sigma_{XO}^{2} \right\}^{\frac{1}{2}}$$
(42)

and the alongwind cloud length (L{x}) for a point source in the layer at the distance x from the source is

$$L\{x\} = \frac{0.28 (\Delta \overline{u}) (x)}{\overline{u}} \qquad . \tag{43}$$

Here,  $\Delta \bar{u}$  is the vertical wind speed shear in the layer and is defined as

$$\Delta \bar{u} = \bar{u}_T - \bar{u}_R , \qquad (44)$$

and  $\sigma_{xo}$  is the standard deviation of the alongwind source dimension in the layer at the point of cloud stabilization. The above equation for L $\{x\}$  is based on the theoretical and empirical results reported by Tyldesley and Wallington (1967) [5], who analyzed ground-level concentration measurements made at distances of 5 to 120 kilometers downwind from instantaneous line-source releases.

In summary, it should be pointed out that the standard deviations of the vertical, crosswind, and alongwind terms represent the cloud dimensions ( $L_i$ ), that is

$$L_{i} = 4.3 \sigma_{i} . \tag{45}$$

The factor 4.3 represents the 97 percent confidence level of a normal distribution. Hence, the initial source dimension is translated into the standard deviation initially for modeling. The standard deviations  $\sigma_x$ ,  $\sigma_y$ ,  $\sigma_z$  give the cloud size during the diffusion process.

## SECTION III. NASA/MSFC SURFACE-LAYER DIFFUSION CALCULATOR PROGRAM

The NASA/MSFC Multilayer Diffusion Calculator Program has been specialized to provide the NASA/MSFC Surface-Layer Diffusion Computer Program in accordance with the discussion given in the last section. The specialization amounts to restricting the applicability to the surface prediction of the transport of HCl for the Titan III. In addition to these modifications, the NASA/MSFC Surface-Layer Diffusion Program also incorporates algorithms to calculate a meteorological profile and cloud rise.

Data requirements are a rawinsonde sounding from the surface to about two kilometers (or about 7000 feet) and the standard deviation of the wind azimuth at the surface. This program utilizes this information to generate graphs for: (1) meteorological profile, (2) the temporal history of the exhaust cloud rise to the point of stabilization, (3) the surface centerline concentration and dosage of HCl along the cloud path, and (4) the surface HCl isopleth along with this path. The operations of the NASA/MSFC Surface-Layer Diffusion Program is straightforward since inputs are called through calculator display. The overall block diagram, showing how the routines are interfaced, is given in Figure 1.

Our initial discussion in this section affords a description of the operation of the program along with its inputs and outputs. The second part gives detailed instructions for operating the NASA/MSFC Surface-Layer Diffusion Program. These instructions are summarized in Appendix A. The last part gives operational sophistication for processing time reduction.

## A. Overviews Of The Computations, Inputs, And Outputs Performed In Each Routine

The NASA/MSFC Surface-Layer Diffusion Program is divided into two parts to meet storage limitations of the HP9820 calculator. The first part is the MET routine and the second part is the diffusion routine (Fig. 1). During operations, the met routine, which generates the inputs for the diffusion model, is erased after cloud rise computations with the loading of the diffusion routine. The meteorological and cloud rise data in storage are not affected during this transition.

Both of these routines can be subdivided into sections where a computational or plotting procedure is carried out. The met routine performs four processes — two involved with raw data processing, the other two with plot generation. These sections are the sort, convert, meteorological profile, and cloud rise. Similarly, the diffusion routine contains four sections of which three sections are involved in generating the two diffusion mappings and the other section contains the basic algorithms for the concentration and dosage calculations. These are the centerline dosage and concentration section, cloud path section, isopleth section, and the FA subroutine. The FA subroutine differs from the other sections in that this subroutine is utilized more than once in part of the program.

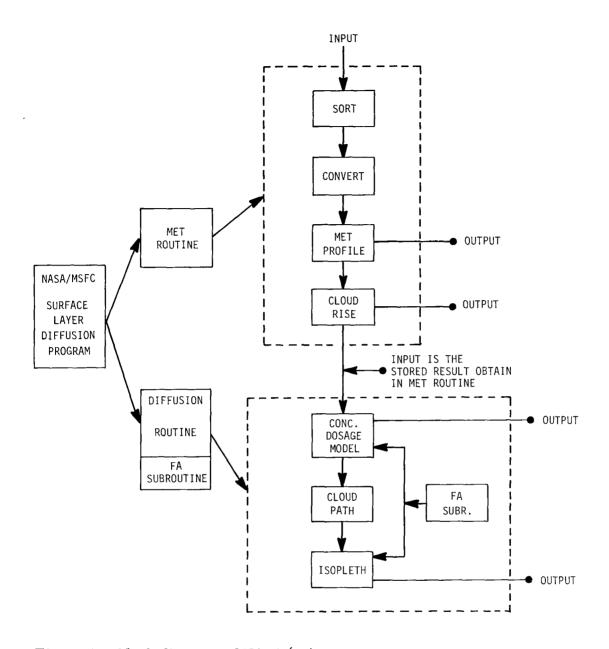


Figure 1. Block diagram of NASA/MSFC Surface-Layer Diffusion Program.

The first section of the met routine, the sort routine, which takes the rawinsonde sounding (Table 1), sorts out the data sets (pressure, temperature, altitude, and wind velocity) by pressure and places these sets in descending order into storage. Termination factors for ending processing loops are derived at the completion of data entry for use in the met routine.

The second section is the convert routine which takes the data from the sort routine and converts it to parameters which are compatible with our calculations. Unit conversions are feet to meters (altitude), knots to meter/seconds (wind speed), and dry bulb temperature to potential temperature. If the altitude for a data set was not given in the rawinsonde sounding, the altitude is calculated by using Equation (4). The dry bulb temperature is plotted in this section (Fig. 2) and options exist at the end of this section to record the processed meteorological data. The data from a formerly processed sounding can be reanalyzed by introducing it at the end of this section.

TABLE 1. INPUT DATA FOR THE SORT ROUTINE

TEST NBR 0475 RAWINSONDE RUN CAPE CANAVERAL MTA. FLA. 12/20/73 1545Z ASCENT NBR 0092

	ALT FT.	WDIR	wkts	ТЕМР	DEW PT	PRESS	RM	AS HUM	DEN	IR	VS	ws
ĺ	16	020	12	17.0	10.8	1022.0	66	9.63	1220.8	330	668	1 1
Ì	1000	033	19	13.5	5.3	986.5	57	6.74	1194.8	307	659	12
1	2000	056	15	11.6	2.5	951.4	58	5.96	1160.2	295	657	8
	3000	031	12	15.1	-16.6	917.6	9	1.25	1108.0	254	661	6
İ	4000	090	10	14.5	-18.1	885.1	8	1.10	1071.2	245	660	2
-	5000	079	5	13.7	-19.4	853.7	8	.99	1035.8	236	660	4
١	6000	052	3	12.1	-20.4	823.2	8	. 91	1004.6	229	658	2
l	7000	057	_4	9.9	-22.4	793.6	8	.77	975.9	222	655	1

I	MANDATORY LEVELS							
ĺ	ALT FT.	WDIR	WKTS	TEMP	DEW PT	PRESS	RM	
	626	043	15	14.8	7.4	1000.0	61	
	2043	057	15	11.8	1.5	950.0	56	
	3539	033	11	14.7	-17.4	900.0	9	
	3120	076	5	13.6	-19.5	850.0	8	
	6751	053	4	10.5	-21.8	800.0	8	

SIGNIFICANT LEVELS				
TEMP	DEW PT	PRESS		
17.0	10.8	1022.0		
13.4	5.2	986.0		
11.2	4.7	955.0		
9.0	-23.5	782.0		

The third section, the meteorological profile, plots out wind direction, wind speed, and potential temperature (Fig. 2). Thus this routine is strictly peripheral to the calculations since it does not contain any data processing.

The last section of the met routine is the cloud rise section, which affords the temporary history of the rise of the exhaust cloud during the initial fifteen minutes of transport (Fig. 3). The kinematic and geometric parameters associated with the stabilization of the exhaust cloud are stored for later operations.

The first section of the diffusion routine is the concentration and dosage section. This generates concentration and dosage levels over a 20 km distance, initiating at the point of cloud stabilization (Fig. 4). Maximum concentration is calculated along with its down range location and plotted in order to give the range of possible concentration isopleths that can be plotted. The second section is the cloud path routine which predicts and plots the exhaust cloud path on a graphical map of the launch area (Fig. 5). This same graph also incorporates the predicted HCl concentration isopleths. Indices with time relations are generated to show the location of exhaust cloud stabilization,

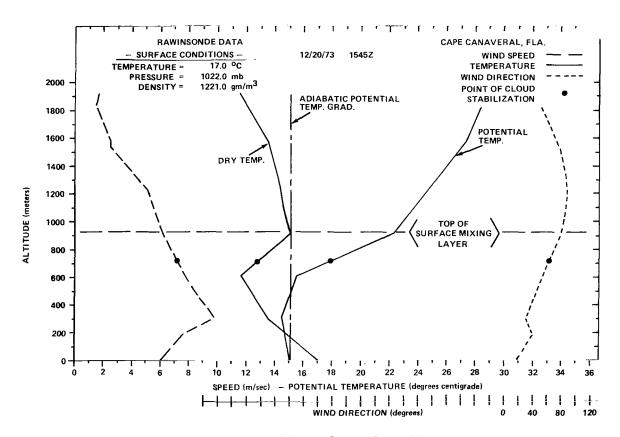


Figure 2. Meteorological profile.

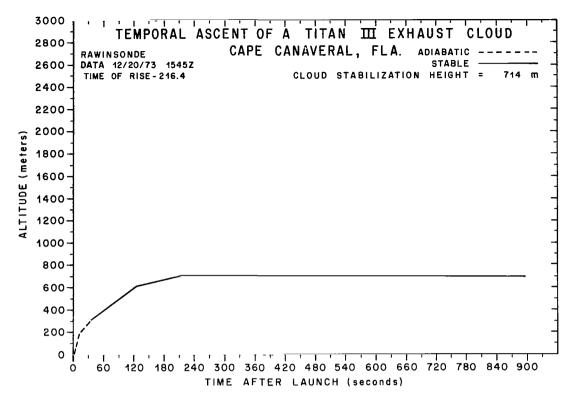


Figure 3. Temporal ascent history of exhaust effluent cloud.

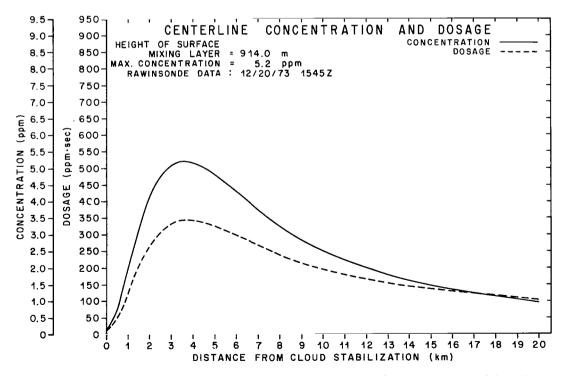


Figure 4. Centerline concentration and dosages for hydrogen chloride.

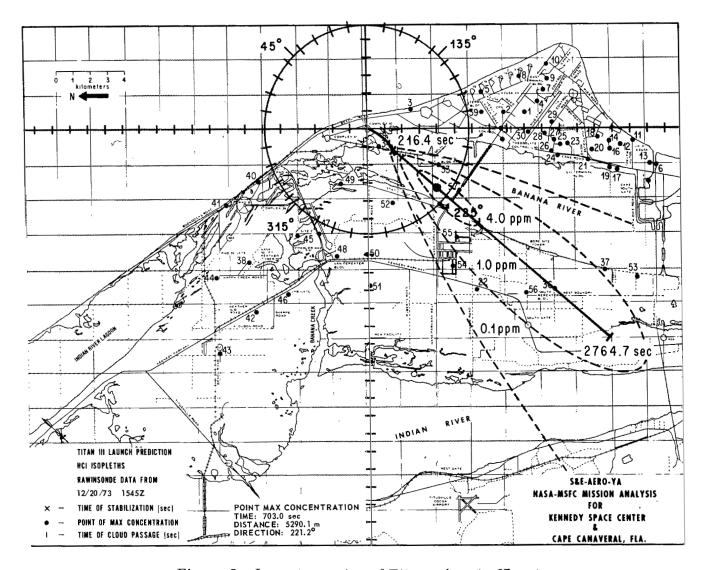


Figure 5. Impact mapping of Titan exhaust effluents.

the point of maximum concentration, and the cloud's arrival time at some arbitrary down range distance. The third section, the isopleth routine, is used strictly for plotting, since it illustrates where contours of constant concentration of HCl reside in the launch area. The routine recycles without termination, permitting several isopleths to be plotted. Termination is not necessary, since this is the last operational portion of the program.

The FA subroutine in the diffusion routine contains portions of the specialized version of model 3 from the NASA/MSFC Multilayer Diffusion Models. Factors produced by this subroutine are combined by other routines to form concentrations and dosage values.

Data requirements for the Surface-Layer Diffusion Calculator Program fall into two categories: (a) the parameters governing the exhaust cloud source as defined by the vehicle and (b) the meteorological parameters of the surface layer obtained from a rawinsonde sounding and surface measurements.

The exhaust cloud source data for the Titan III is programmed into the command lines of the program as a portion of a constant to its respective equation. In Titan III case studies, reentry of these parameters is not necessary. To alter source data requires the reprogramming of command lines in accordance with the discussion in Section II.

The four parameters which make up the source data are: (a) the heat released by the Titan III exhaust ( $Q_M$ : Equation 12) and utilized in the buoyancy term ( $F_I$ : Equation 11). (This is located in line 60 of the met routine); (b) the fractional amount of material expelled into the surface layer (f) and required for calculating the mass source strength ( $Q_M$ : Equation 21). (This is located in line 0 of the diffusion routine); (c) the entrainment coefficient ( $\gamma$ : Equation 7, 8), utilized in the cloud rise relations and in the calculations for the standards in wind shear, ( $\sigma_X$ ,  $\sigma_Y$ ,  $\sigma_Z$ ): Equation 45). (This is located in line 60 of the met routine and in line 1 of the diffusion routine); and (d) the molecular weight of the toxin of interest (M), and required for calculating  $Q_M$ . (This is located in line 1 of the cloud routine). Respective values for  $Q_H$ , f,  $\gamma$ , and M for the Titan III C in HCl measurements are 8.42915 × 10<sup>6</sup>, 1.54486 × 10<sup>9</sup>, 0.64, and 36.5 [4].

Meteorology data is entered into the HP9820 computer through the keyboard, on display calls from the program. The program offers data card loading and recording to facilitate operations in case study reproductions and in forming data references. The five meteorology profile parameters utilized by the surface layer diffusion routine are altitude with its respective wind direction, wind speed, temperature, and pressure. Units are either English or metric, since the necessary conversion routines are contained within the program. Data points for these parameters are within the surface layer, which is assumed to be between ground level and 2.1 km. These five meteorology parameters constitute one data set. Nineteen data sets are the maximum number which storage can handle, although at least fifteen are required to prevent processing overflows. Generally, from fifteen to eighteen sets of data can be derived from one rawinsonde recording. The date of the rawinsonde data is requested by the program to be used for referencing the outputs. The input occurs following meteorology profile parameter plots.

The input for the diffusion routine correspond to wind direction, wind speed, altitude, and standard deviation of wind azimuth at the top of the layer (when available), and standard deviation of wind azimuth at ground level. These parameters are selected from the meteorology profile to indicate where the top of the surface layer mixing is. The last input appears in the center of the isopleth routine. This requires a concentration value with which its corresponding isopleth is to be plotted.

The output generated by the NASA/MSFC Surface-Layer Diffusion Calculator Program are designed to afford both a spatial and temporal prediction for the foot print from the transport of the solid rocket effluents. In addition, values which are important in the analysis of the resulting predictions are printed on these graphs. To reduce the data reduction time, separate programs exist to format each graph with respect to setting up the proper coordinate system and the proper labelings (in normal operation, we have found that reproducing these from the master is a time saver).

The meteorological profile represents the atmospheric kinematics and theromdynamics which are relevant to the dispersion of the rocket exhaust effluents. The potential temperature is given in this profile because it is the primary determining parameter in predicting the altitude at which the rocket exhaust cloud will stabilize. The selection of whether the adiabatic or the stabilized cloud rise relation is utilized in the calculation of the cloud stabili-

zation height depends on the adiabatic lapse rate relative to the surface. The adiabatic potential temperature regime, the stable potential temperature regime, and the stable potential temperature are defined on the meteorological profile. The dynamic range for wind speed and temperature (potential and dry bulb) profiles is from 0 to 36.5 m/sec and from -4°C to 36.5°C, respectively. Wind direction scaling varies in increments of 20 degrees to actual conditions with a maximum range of 0 to 360 degrees. The program calls for a stop after each profile is plotted so that the color of the pen can be changed, thereby making each profile more distinguished from the other profiles. In addition to the profiles, the surface conditions (such as temperature, pressure, and density), the date, and time of the rawinsonde soundings are printed on the meteorological profile.

The cloud rise graph affords the temporal history over the initial fifteen minutes after rocket ignition. The type of atmosphere (adiabatic or stable) is indicated on the delineation of the cloud rise. In addition, the altitude and transport time to cloud stabilization are printed out.

The third graph routine is the dosage and concentrations of HCl along the centerline of the exhaust cloud path. This delineation ranges over the first 20 km after cloud stabilization. The height of the surface mixing layer and the maximum concentration of the HCl are printed out. The concentration and dosages of other rocket exhaust constituents can be obtained by multiplying by the constants of proportionality which are:

```
    1.73 for CO (ppm)
    0.11 for CO<sub>2</sub> (ppm)
    2.22 for Al<sub>2</sub>O<sub>3</sub> (mg/m<sup>3</sup>).
```

The maximum values of the concentration and dosage which can be drawn are 9.5 ppm and 950 ppm-sec.

The final output is a graph of the HCl isopleths of constant concentration. This routine plots the actual cloud path showing where the cloud stabilization and the maximum concentration occur. In addition, an arbitrary point is marked with the time of cloud passage to enable the user to determine the passage time at any point (note that we assume that the cloud has a constant velocity after cloud stabilization). Numerical printouts give the time of cloud passage, the

distance from the launch site, and the angle to the point of maximum concentration.

The details of the operation of this calculator program in the next section will help to illuminate points which are not made clear here.

## B. Operating Instructions For The NASA/MSFC Surface-Layer Diffusion Calculator Program

A detailed set of operating instructions for the NASA/MSFC Surface-Layer Diffusion Program is provided in this section. While this set of instructions is helpful when initially running the program, the summary of operating instructions given in Appendix A should be sufficient in normal operations. The NASA/MSFC Surface-Layer Diffusion Program is loaded into the machine (HP9820) in two parts, where the second part depends on the first part. Each part has two sections, each affording one of the four graphs. The normal operation of this program can be summarized as loading the met routine and then running the program. Data will be called for in the display. The user must terminate the data entry with a GO TO 12. Otherwise all other instructions are called for until the cloud rise graph is completed. Loading of the second part (the diffusion routine) requires a bit more care, since a subroutine must also be entered. Once this is completed, the remainder of the program is straightforward.

In the following instructions, the names of calculator keys are underlined and the displays are both underlined and set off by quotation marks.

#### Part I. MET Routine

#### Section I. METEOROLOGY PROFILE

#### A. Loading Met Routine

- 1. The met routine consist of two cards, 4 sides, numerically listed for loading.
- 2. Press ERASE MEMORY, LOAD, EXECUTE. Insert each card in numerical order into the card reader. After each side is loaded, 'NOTE 14' will be displayed, with exception of side four. Press EXECUTE and insert the next side.
- 3. If side four is followed by a 'NOTE 14" display, the calculator has misread program cards. Press ERASE MEMORY, and reload the met routine.
- 4. Place graph on plotter for Meteorological Profile (use when reprocessing data).

#### B. Data Cards

Input data can be recorded in converted form on magnetic cards for reference to a particular case study. Magnetic cards may also be used to enter data into the calculator. These operations center around the 'LOD OPT' display and will be discussed later.

- C. Entering Data Through The Keyboard
- 1. During processing the calculator will idle, generating a display for data entry.
- 2. Following the display, type in the corresponding data and press RUN PROGRAM.

Example

"TIME" is displayed.

Enter the time through the keyboard and press RUN PROGRAM.

- D. Initiating the Met Routine Loading Data by Cards
- 1. Press RUN PROGRAM until 'Z' is displayed. Press STP, GTO 26, RUN PROGRAM. This jumps over data input and conversion routines required for raw information. 'DS' will be displayed. Surface density is stored on the magnetic card so data entry for this call is unnecessary. Press RUN PROGRAM.
- 2. The card drive will activate, but in the record mode for converted raw data. Press STP, RUN PROGRAM.
- 3. "LOD OPT" will be displayed. Press LOD, EXECUTE, and load data into the calculator.
  - 4. Dry bulb temperature will not plot out with the use of data cards.
- E. Initiating the Met Routine With Raw Data
- 1. Press RUN PROGRAM until "Z" is displayed. Enter the following rawinsonde data to the corresponding display in accord with C.

Z - Altitude (ft. or meters)  $\theta - Wind Direction$  (Deg) U - Wind Speed (knot or meters) T - Temperature (Dry Bulb °C) P - Pressure (mb)

Each group of five parameters constitutes one data set. Fifteen data sets are required for processing, although nineteen sets are the maximum to be stored. In cases of missing data points, the following may be substituted.

Missing Parameter	Substitute
Altitude	0
Wind Direction	1000
Wind Speed	-1

- 2. If data was miskeyed into the calculator, see the next part for data correction or press STP, GTO 0, RUN PROGRAM and reenter data sets.
- 3. After storing the rawinsonde data, press <u>STP</u>, <u>GTO 12</u>, <u>RUN PROGRAM</u>. "METRIC?" will be displayed. If rawinsonde data were in metric units, press <u>GTO 19</u>, <u>RUN PROGRAM</u>, to jump over the conversion routine, otherwise press <u>RUN PROGRAM</u>.
  - 4. Dry bulb temperature will plot out. Change pen colors.
  - 5. ''DS'' will be displayed. Enter in the surface density in  $gr/m^3$ .
- 6. The card drive will activate. Converted data can be recorded by inserting a data card into the card drive; otherwise, press <u>STP</u>, <u>RUN</u> <u>PROGRAM</u> (when initally using the program, it is recommended that data be recorded). A 'NOTE 14'' display afterwards indicates a data error or a recorder malfunction. In such case, press <u>STP</u>, <u>EXECUTE</u>, <u>GTO 26</u>, <u>RUN PROGRAM</u>. With the 'DS'' display press RUN PROGRAM and record again the data. After the second appearance of 'NOTE 14'', if any, see the next part for data correction.

7. "LOD OPT" will be displayed. This concerns only data card loading and it should not be associated with raw data.

#### F. Plotting the Meteorology Profile

- 1. After the "LOD OPT" display, either with keyboard or card loaded data, press RUN PROGRAM.
  - 2. Wind speed will plot out. Change pen colors and press RUN PROGRAM.
- 3. Potential temperature and the adiabatic potential temperature gradient will plot out. Change pen colors and press RUN PROGRAM.
  - 4. Wind direction and its scaling will plot out.
- 5. ''MONTH'' will be displayed. Enter in the date of the rawinsonde data with the following displays: 'MONTH', 'DAY', 'YEAR', 'TIME''.
- 6. Date of rawinsonde data and surface temperature, pressure, and density will plot out.
  - 7. The ''C-GRAPH'' displays the end of the meteorology profile routine.

#### Section II. CLOUD RISE

- 1. Set up the plotter. After the 'C-GRAPH' display, press RUN PROGRAM. If an adiabatic condition exists, plotting will commence until the 900 second limit is reached or until a stable condition prevails.
- 2. Once processing has stopped or plotting did not initiate after pressing RUN PROGRAM at the start, change pen colors and press RUN PROGRAM until the completion of the plot. For each time the program stops, a numerical readout will be displayed, which is the time factor calculated for given altitude. This effect indicates the processing of a stable condition.
- 3. After completion of the printing, 'Stable Atmosphere' will be printed out beneath the line segments which it represents. Here the program will stop for a pen color change. Press RUN PROGRAM to resume processing. Dates of rawinsonde data, time to maximum cloud rise, and 'adiabatic atmosphere' (if one exists) will be plotted out. This completes the graph and the met routine.

4. If a note statement appears or the cloud rise plot seems unrealistic in accordance with the meteorology profile, data is in error. Most frequently trouble will arise in registers R1, the surface temperature in kelvin; R2, surface pressure in mb; and R3, surface density in gr/m³. See the next part for data correction and Part VII for recycling the program.

#### Part II. Diffusion Routine

#### Section III. CENTERLINE CONCENTRATION AND DOSAGE

- 1. The cloud diffusion routine requires two cards. Three sides are used for the concentration and dosage routines and one side for the FA subroutines. MEMORY ERASE should not be depressed, since data stored in the previous routine are used to begin loading. Press DEF FA, EXECUTE, LOD, EXECUTE and insert the FA subroutine program card. This is stored into the FA register and may be called back for inspection by pressing DEF FA, LIST or DEF FA, RECALL. To load in the cloud routine, press END, EXECUTE, and insert the cloud routine program cards. The END function separates the program mode from the subroutine mode.
- 2. Set up the plotter for the centerline graph. Press <u>RUN PROGRAM</u> until ' $\theta$ T'' appears in the display. Select from the meteorology profile the altitude, wind speed, and wind direction for the top of the surface layer. With each of the following displays, enter its corresponding data:

 $\theta$ T-wind direction at the top of the layer,

UT-wind speed at the top of the layer,

 $\theta$ AR-wind deviation at ground level,

 $\theta$ ATK-wind deviation at the top of the layer, and

HL-height of the layer.

3. After the last data input, plotting will initiate. If, for some period of time the calculator should idle with no apparent output, storage data is in error. Go to the next part for data correction and for recycling the program.

4. Following the concentration and dosage plot, the height of the surface layer, maximum concentration, and the date of the rawinsonde data will be printed out on the graph. The ''ISO-GRAPH'' display indicates the end of this part of the cloud routine.

#### Section IV. ISOPLETH

- 1. When ''ISO-GRAPH'' is displayed, set up the plotter for the isopleth graph, change pen colors and press RUN PROGRAM.
- 2. Cloud path will plot out in a relatively straight line, with the majority of its deviations centered around the origin. If there are any signs of irregularities similar to a spike or a jump in the line, miscalculated data was stored in memory. To continue with processing will lead to inaccuracies. For data correction and to recycle the program, see the next part.
- 3. In many cases the cloud path will go off scale and the arbitrary down range time value will be missing. This is a processing error. No action is necessary since this has no effect on the output. To prevent this, press STP before the cloud path reaches the limits of the graph, then press GTO 25, RUN PROGRAM.
- 4. Following the cloud path plot, time of cloud rise, an arbitrary down range time factor, and the position of maximum concentration are marked out. Maximum concentration parameter and date of the rawinsonde data are then plotted out.
- 5. ''C'' will be displayed. Enter in a concentration greater than 0.1, for practicability, and no less than maximum concentration.
- 6. The time before initial plotting is dependent on the proximity of the entered value to maximum concentration. This lengthens as the two values approach each other. In some cases, a note statement will be displayed, indicating that the calculator is unable to determine the initial plotting location. Enter a new concentration value by pressing STP, EXECUTE, END, EXECUTE, GTO 31, RUN PROGRAM. If only one point is plotted and the calculator idles for a period of time, enter in a new concentration value.
- 7. During plotting, if the isopleth should go off scale, the program will recycle for the entry of a new concentration value. If the isopleth remains within the limits of the graph, 'NOTE 2" will usually appear. Press RUN PROGRAM to complete the isopleth. If 'NOTE 2" appears a second time, press STP, EXECUTE, GTO 31, RUN PROGRAM."C" will be displayed.

- 8. To initialize an isopleth with a concentration value, press  $\underline{\text{SET/CLEAR}}$  FLAG during its plot.
  - 9. This completes the isopleth and the cloud routine.

The next part affords some helpful tips on how to correct errors and to short cut data processing steps.

#### C. Corrective Procedures And Operational Shortcuts

Since errors can be introduced into the final results by incorrect inputs to the NASA/MSFC Surface-Layer Diffusion Program, a careful check should be made of each graph for obvious errors - especially, the meteorological profile. If an error is detected, it is time saving to be able to correct the error and continue processing without having to go to the beginning and start the program at step 1. This part discusses the potential errors and how to correct them. Then the necessary steps to recycle the program are given.

Since there may be times when only the output of a specific section is desired, the methods for by-passing unwanted output are also considered. Appendix B gives the location of the meteorological, cloud rise, and diffusion algorithms used in the program along with a flow chart and listing of the program.

The need for a data correction can be indicated by the appearance of a note statement (not normal to the operation), or the results of an obviously questionable output which are the indications of erroneous data in storage. Mistakes in keying data into the calculator are generally the source of the error; although rarely, a misread in the program card would lead to a miscalculation. The case run can be salvaged if the error is located and corrected.

- With the display of "NOTE A in B", where A is the note type and B is the line number where the error takes place, one should identify the note and its possible cause.
  - NOTE 1: Illegal instruction. Program card misread.
  - NOTE 2: Improper store or taking the square root of a negative number.
  - NOTE 3: Illegal instruction. Program card misread.

- NOTE 4: Illegal instruction. Program card misread.
- NOTE 5: Miscount in data increment calling for a nonexisting or program register.
- NOTE 6: Miscount in data increment storing a value in a nonexisting or program register.
- NOTE 8: Recalling nonexisting program line.
- NOTE 9: Illegal instruction. Program card misread.
- NOTE 10: Underflow or overflow.
- NOTE 12: Illegal instruction. Program card misread.
- NOTE 14: Magnetic card operation incomplete.
- NOTE 16: Printer out of paper.
- To correct program card misreads, press LIST, following the note statement. Inspect the printout with the listings given in Appendix B. Once the invalid line has been located, press GTO B (B is the program line), RECALL, BACK, and press DELETE until only the line number remains (ir. ''64:''). Type in the corresponding command from Appendix A and press STORE (see Part VII to recycle the routine). If the program card misread is extensive, press ERASE MEMORY and restart the program.
- For a single invalid data point, press GTO B, where B is the line number from the note statement, and LIST. Allow a few lines to print out and press STP.
- The first line in the listing contains some register (R0, R1...R103, A, B, C, X, Y, Z) and the operation where the error developed. The content of these registers can be determined, for example, by pressing B, EXECUTE, where B is the register of interest.
- To correct data, press  $N\rightarrow B$ , EXECUTE, where N is the new value for register B.

- If invalid data cannot be located by the above method or there are a number of data points at fault, a data dump will be required. To initiate, press  $0 \rightarrow X$ , EXECUTE, PRT RX: JMP (X+1 $\rightarrow X$ ) 100, and press EXECUTE until about a hundred figures have been printed out. R0 on down consecutively to R103 will be listed and in this way irregularities in groups of data may be found. To insert correction press  $N \rightarrow B$ , EXECUTE where N is the new value for register B.
- Following data correction, the program should be recycled back to the beginning of the routine (Part VII). In cases where data correction cannot be handled, press ERASE MEMORY and restart the program.

Portions of the program, where an output is not needed, may be jumped over during processing if the data required is accessible. The meteorology profile, cloud rise, and isopleth are a waste in operating time if the only output required is the centerline concentration and dosage. The method described below will show how this may be done.

#### a. Cloud Rise Plot

- (1) To plot cloud rise, the met routine should be loaded into the calculator and the converted rawinsonde data must be present. If data is absent from storage, it should be entered by magnetic card at the 'LOD OPT'. Displayed or keyed in at the beginning of the met routine and followed through until 'LOD OPT' appears.
- (2) After the "LOD OPT" display, press GTO 58, RUN PROGRAM. Cloud rise will plot out but the date of rawinsonde data will be in error, unless Month R90, Day R91, Year R92, Time R93 is entered.
  - b. Centerline Concentration and Dosage
    - (1) Load the cloud routine in the following sequence

#### PRESS MEMORY ERASE

PRESS <u>DEF</u> <u>FA</u>, <u>EXECUTE</u>, <u>LOD</u>, <u>EXECUTE</u>. Insert FA subroutine.

PRESS END, EXECUTE, LOD, EXECUTE. Insert cloud routine.

(2) Rawindsonde data is not required, but the following is, although it is not called for by the program.

$\mathbf{DATA}$	REGISTOR
Surface Temperature °K	R1
Surface Pressure mb	R2
Time of Maximum Cloud Rise	R85
Maximum Cloud Height	R86

ENTER uncalled data by pressing  $N\rightarrow RB$ , EXECUTE where N is the data and B is the corresponding register number.

- (3) Press RUN PROGRAM until  $^{1}\theta T^{1}$  is displayed. Continue operations with instructions in Part III.
  - (4) Date of rawinsonde data will be in error.

#### c. Isopleth

- (1) Isopleth can only be generated by running the entire program. Data requirements are too large to calculate and enter. There are, however, methods to reduce processing time.
- (2) Load in the met routine and input data by card or keyboard. Follow the meteorology profile instructions, Part I, until the "LOD OPT" stage ends. Press GTO, 58, RUN PROGRAM.
  - (3) Let the cloud rise routine run out completely.
- (4) Follow through with instructions in Part III for centerline concentrations and dosage.
- (5) After the plot has passed point of maximum concentration, press END, EXECUTE, STP, EXECUTE, GTO 13, RUN PROGRAM.
  - (6) Follow the instructions for isopleth, Part IV.

Recycling sections of the program, following a data correction, is required to reset calculations with the altered data values. Simple jumps from the point of correction to the start of a routine eliminate the need to run several sections of the program to generate an output. The commands listed below may be used in any point of a given routine for recycling back to its beginning unless stated otherwise.

#### a. Recycling Meteorology Profile

- (1) While entering raw data, if a miskeyed data point was inputted, two methods may be taken for correction. Either search for the data point and correct it, as in Part V and continue to enter data sets, or wipe out stored data and reenter the data sets by pressing STP, GTO 0, RUN PROGRAM.
- (2) If more than 19 data sets were entered, command lines are erased, rendering the program and data unsalvable. Press <u>ERASE</u> MEMORY and reload the met routine.
- (3) To recycle during the plotting sequence, press STP, 4-X, EXECUTE, GTO 27, RUN PROGRAM. If the graph is to be replotted, dry bulb temperature will be missing.

#### b. Recycling Cloud Rise

Processing should be in the stop mode. Press R1-273.15→R1, EXECUTE, GTO 58, RUN PROGRAM.

#### c. Recycling Centerline Concentration and Dosage

Press STP, EXECUTE, END, EXECUTE, GTO 0, if new top of the layer parameters are to be entered, otherwise GTO 2, RUN PROGRAM.

#### d. Recycling Isopleth

- (1) In the cloud path mode press R97→R64, EXECUTE, GTO 13, RUN PROGRAM.
- (2) In the isopleth mode, press STP, EXECUTE, END, EXECUTE, GTO 13, RUN PROGRAM.

While these corrective proceedures cover most cases that will be encountered, there will be times when the operator of the program will want to employ his own techniques. While these proceedures may normally be satisfactory, great care should be exercised due to the complexity of the computational interfaces.

#### SECTION IV. CONCLUSION

The NASA/MSFC Surface-Layer Diffusion Calculator Program affords a real-time, on-line transport description for the environmental dispersion of the Titan III rocket exhaust effluents utilizing ellipsoidal dispersion in a Eulerian reference frame, as given by Model 3 of the NASA/MSFC Multilayer Diffusion Model. This calculator program has been successfully employed at Cape Kennedy to make real-time predictions of the transport of Titan III effluents for the deployment and evaluation of the Langley Research Center's effluent monitoring network.

The primary difference between the NASA/MSFC Surface-Layer Diffusion Calculator Program and the NASA/MSFC Multilayer Diffusion Computer Program is that the calculator program has been specialized for a specific vehicle, the Titan III, and a specific altitude, the surface layer. The calculator program incorporates the meteorological and cloud rise algorithms into one program, permitting a rawinsonde sounding to be the input; whereas, the computer program utilizes the results of these algorithms. The predictions obtained from both programs are identical for the surface effects from these effluents. The data preparation and reduction time for the calculator program is about a half hour, as compared to two hours for the computer program. The computer program, however, does afford a more detailed prediction than can be obtained with the calculator program.

The NASA/MSFC Surface-Layer Calculator Program has features which make it a desirable vehicle for use in making diffusion predictions for the Titan III launches, especially where the graphical output satisfies the prediction required. The NASA/MSFC Multilayer Diffusion Computer Program is a better vehicle for use in general analysis work where flexibility and detail in the results are required. Thus, while the application of the calculator program to diffusion analysis is a step forward in on-line, real-time predictions, it does not eliminate the need for the computer program in diffusion analysis.

# APPENDIX A. SUMMARY OF THE INSTRUCTIONS FOR THE OPERATION OF THE NASA/MSFC SURFACE-LAYER DIFFUSION CALCULATOR PROGRAM

The objective of this summary is to provide a guide in the operation of the calculator program. The details for the initial operation of this program are given in Part B of Section IV. Since, in normal operations of this calculator program, only a quick reminder of some key operation is all that the operator will require, the following summary is provided. Even this may be too detailed for many and, therefore, a flow chart for the operation of this calculator program has also been provided (Fig. 6).

The user is reminded that the procedures for treating program errors and eliminating unwanted parts of the program are given in Part C of Section III. This discussion does not include the points covered in this section.

#### Operation of Program

Data called for by the machine is entered in the following sequence.

''<u>DISPLAY</u>'' The program makes the call for data type in data and press RUN PROGRAM.

Met. Profile PART 1. METEOROLOGY PROFILE MET ROUTINE

Load Program 1. This routine consists of two cards, 4 sides, Press MEMORY ERASE and load in the routine.

Loading Data by Card

2a. If data is to be entered by magnetic cards, press RUN PROGRAM until "Z" appears in the display. Press STP, GTO 26, RUN PROGRAM. "DS" will appear in the display though data entry for it is not necessary. Press RUN PROGRAM. The card reader will activate in the record mode for data records, which at this point does not exist. Press STP, RUN PROGRAM. "LOD OPT" will be displayed. Press LOD, EXECUTE, and insert data card. Dry bulb temperature will not be plotted with data card usage. Press RUN PROGRAM to resume processing.

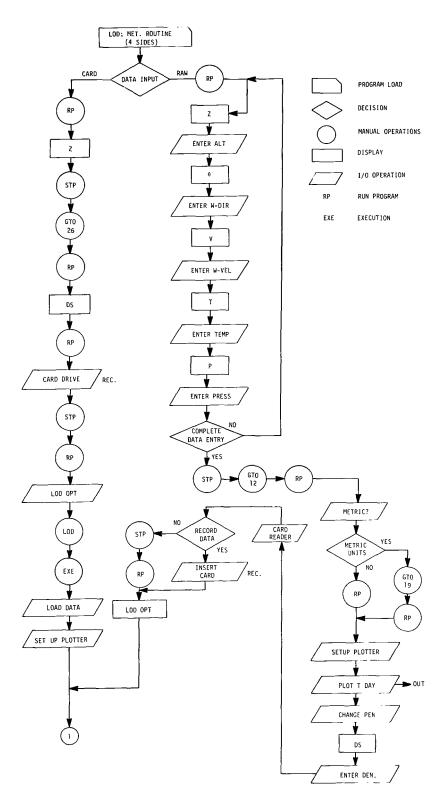


Figure 6. Flow diagram for program operations.

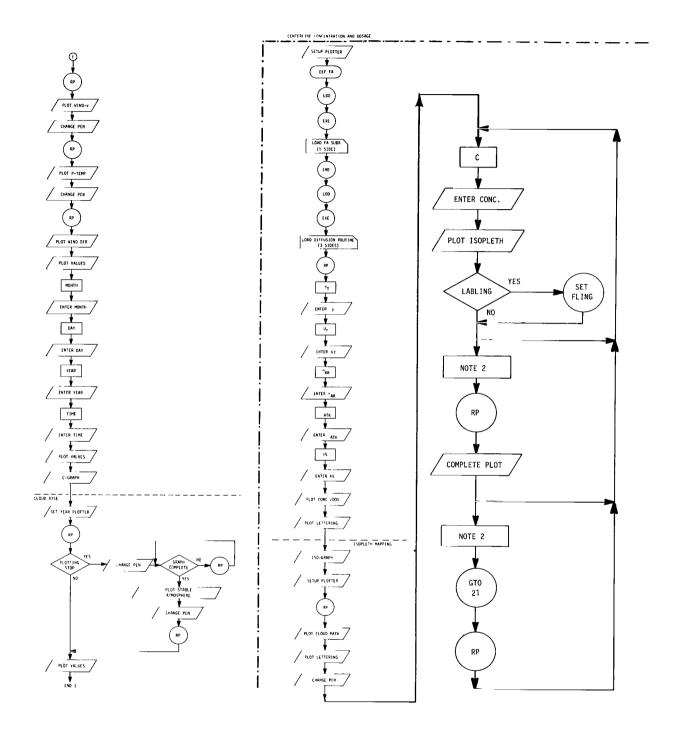


Figure 6. (Concluded).

#### Entering Raw Data

- 2b. The met routine should be loaded. Press RUN PRO-GRAM until "Z" appears in the display. Enter data with the corresponding display. "Z" altitude, "0" wind direction, "V" wind speed, "I" temperature, and "P" pressure. No more than 95 values may be entered. For missing data, the following may be substituted: Altitude: 0, wind direction: 1000, wind speed: -1.
- 3. After entering the last data value, press STP, GTO 12, RUN PROGRAM. 'METRIC?" will be displayed. If data was in metric units, press GTO 19, RUN PROGRAM, otherwise press RUN PROGRAM.
- 4. Dry bulb temperature will plot out. ''DS'' will appear in the display. Set up plotter. Enter surface density value, change pen color, and press 'RUN PROGRAM''.

#### Recording Data

- 5. The card drive will activate. If converted data is to be recorded, insert a magnetic card, otherwise press STP, RUN PROGRAM.
- 6. 'LOD OPT' will be displayed. Press RUN PROGRAM.
- 7. Wind speed will be plotted out, stopping afterwards for a pen color change. Press RUN PROGRAM.
- 8. Potential temperature will be plotted out, stopping afterwards for a pen color change. Press RUN PROGRAM.
- 9. Wind direction, wind direction scale, surface conditions, and date of meteorology data will be plotted out. ''C-GRAPH'' will be displayed.

#### Cloud Rise

#### PART II. CLOUD RISE MET ROUTINE

- 1. Center graph paper on plotter and press RUN PRO-GRAM.
- 2. During processing, if the plotter should stop, change pen color and press RUN PROGRAM as many times as needed to complete the graph.

3. After the plotting of 'STABLE ATMOSPHERE", change pen color and press RUN PROGRAM. This completes the met routine.

#### Centerline Concentration and Dosage Loading Cloud Routine

### PART III. CENTERLINE CONCENTRATION AND DOSAGE CLOUD ROUTINE

- 1. Load cloud routine, 2 cards, 3 sides, and FA subroutine 1 card, 1 side in the following sequence. Do
  not press <u>MEMORY ERASE</u>, press <u>END</u>, <u>EXECUTE</u>,
  LOD, EXECUTE. Load cloud routine.
- 2. Press RUN PROGRAM until 'OT'' is displayed. Enter the corresponding data to each display. OT-wind direction at the top of the layer. OATK-wind deviation at the top of the layer UT-wind speed at the top of the layer HL-height of layerOAR-wind deviation at reference.
- 3. After the final data entry, the program will plot concentration and dosage, and then print out height of layer, maximum concentration and date of meteorology data. This completes the concentration and dosage graph. An 'ISO-GRAPH' will be displayed.

#### Isopleth

#### PART IV. ISOPLETH CLOUD ROUTINE

- 1. Center graph paper on plotter and press cloud routine RUN PROGRAM.
- 2. Cloud path, maximum concentration parameters, and date of meteorology data will be plotted out.
- 3. ''C'' will appear in the display. Enter in a concentration value greather than 0.1 but less than maximum concentration. The corresponding isopleth will be plotted.
- 4. To initialize the isopleth during plotting with a concentration value, press SET/CLEAR FLAG.

- 5. If 'NOTE 2'' is displayed, press RUN PROGRAM to complete the plot. After the second appearance of 'NOTE 2'', press GTO 31, RUN PROGRAM. 'C'' will appear in the display.
- 6. If 'NOTE 2" did not appear, the program cycles back to plot another isopleth, displaying 'C" for a new concentration value.
- 7. This completes the cloud routine. For completion of each graph there is a corresponding routine for axis and lettering.

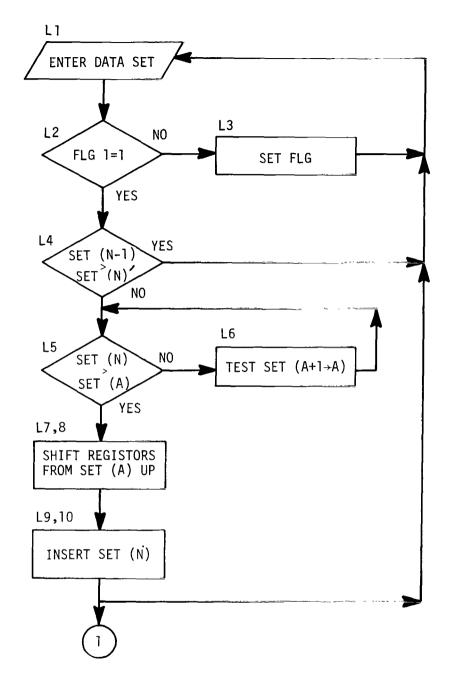
## APPENDIX B. THE NASA/MSFC SURFACE-LAYER DIFFUSION CALCULATOR PROGRAM LISTING AND ALGORITHMS

A listing of the NASA/MSFC Surface-Layer Diffusion Calculator Program, along with a flow chart is given here. In addition, the location of the algorithms used in the program are given.

This calculator program was designed for use on the HP9820 desk calculator, with option 001 (429 total registers). The User Definable Function I-III ROM are required in this program. The program requires all the storage available in this configuration. A listing of the MET routine is given in Table 3, and the Diffusion Routine is given in Table 4. Figures 7 through 14 are the flow diagrams for the program.

Since experience in making diffusion predictions generally leads to a desire to update the algorithms utilized, the meteorological, cloud rise, and diffusion algorithms are given in Table 5. However, the algorithms employed in this program have evolved from years of empirical experience in monitoring the dispersion of the Titan III effluents and, therefore, should require practically no changes.

To assist the user in making his initial predictions with this program and to permit the checking of its operation, we have included some average values for the parameters in Table 6. Table 7 provides the four programs for formating and labeling each of the graphs. In general, time can be saved by making a master for each graph and reproducing it for use with the diffusion program.



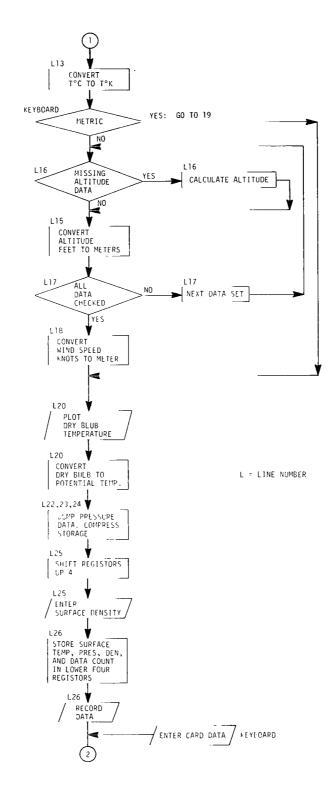
L = LINE NUMBER

SET (N) = THE ENTERED DATA SET

SET (N-1) = LAST DATA SET STORED

SET (A) = DATA SET, INITIATING AT THE FIRST DATA SET

Figure 7. Flow diagram for sort routine.



\_\_\_\_\_

Figure 8. Flow diagram for convert routine.

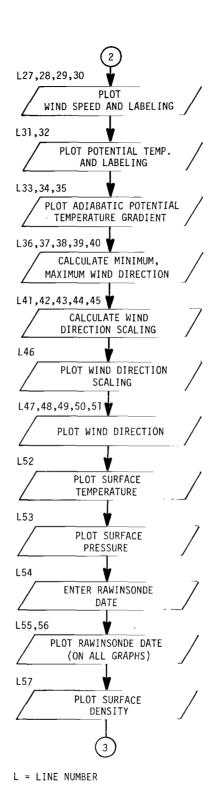


Figure 9. Flow diagram for meteorology profile routine.

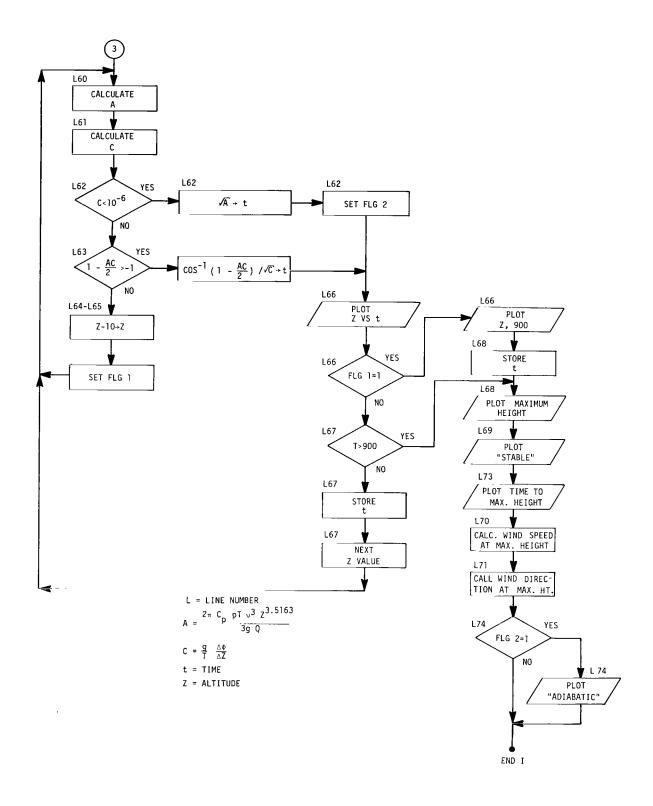


Figure 10. Flow diagram for cloud rise routine.

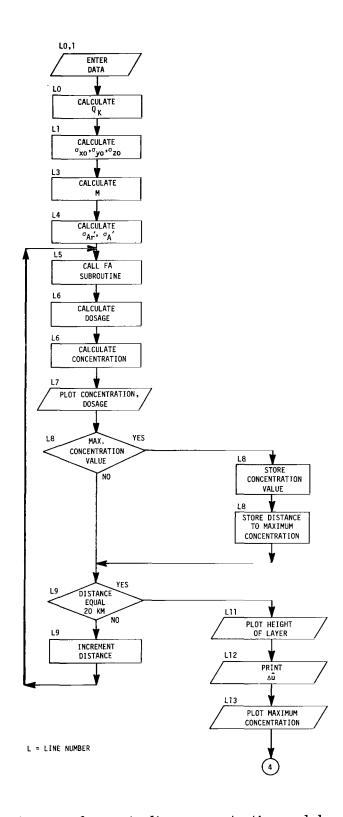


Figure 11. Flow diagram for centerline concentration and dosage routine.

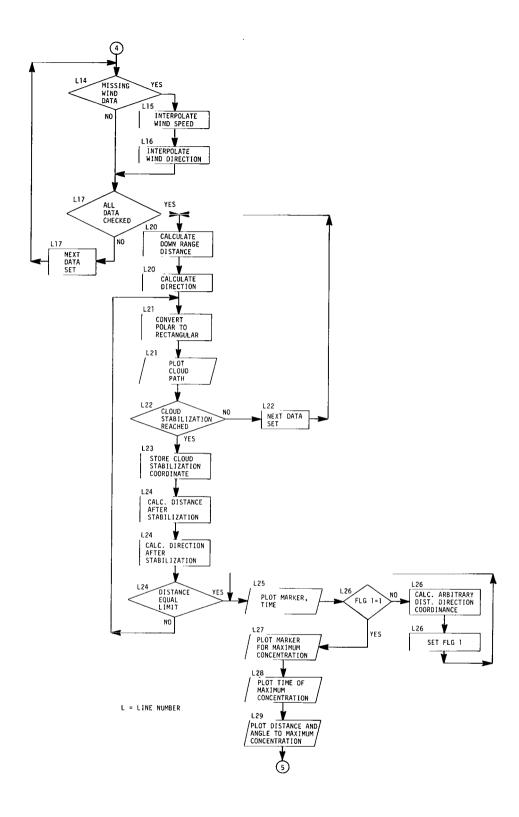


Figure 12. Flow diagram for cloud path routine.

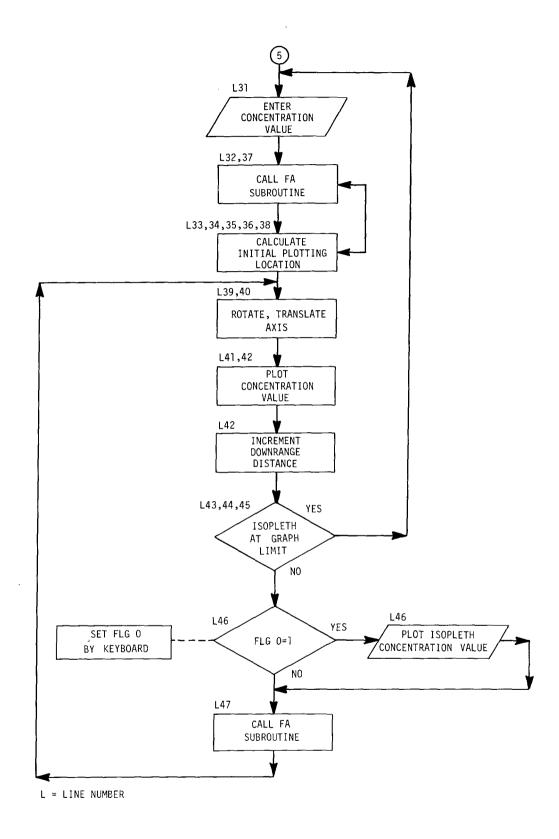
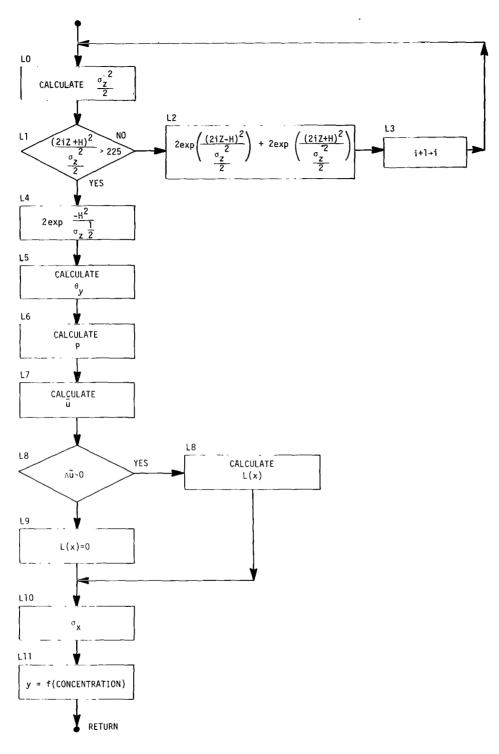


Figure 13. Flow diagram for isopleth routine.



L = LINE NUMBER

Figure 14. Flow diagram for FA subroutine.

## TABLE 2. LISTING OF MET ROUTINE

0: Ø⇒[... Ø⇒X, Ø⇒C; CFG 2; SCL -4, 36.5, -4 00,2400F 1: ĒNT "Z",R(B),"0" \*F(B+1) \* "V" \*R(B+ 2), "T", R(B+3), "P ",R(B+4) 2: IF FLG 2=1;GT0 4 3: SEG 2:5⇒B:GTO 1E 4: IF R(B-1)>R(B+4) \$GT0 11H IF R(B+4)>R(C+4) ;GT0 7H 6: C+5-0;GTO 5H C+748-C+4+ZH 8: R(B+4-X)→R(B+9-X );JMP (X+1+X)>ZH 9: 0+XF 10: R(B+9-X) ≏R(Y+4-X );JMP (X+1→X)>4F 11: B+5+B;0+X;0+C; GTO 1H 12: 3+M;B-5+B;R3+R99 £R4→R98E 13: RX+273.15⇒RX; JMP (X+5+X)>BF 1.1: 0 % CDSP "METRIC ?'∶stP ⊢ 15: .3048RY+RYH 163 IF RY=0:R(Y-5)+2  $9.3 \cdot (R(Y+3) + R(Y-$ 211 21LN (R(Y-1) .R(Y+4)) +RYF 17: IF B: Y; 2+X; Y+5+Y 16TU 15H .515PM→RX;JMP +X +5+X1,B+5H 19: 3→X;R28-273.15→R 97H

20: PLT RX-273.15,R( X-3); RX(1000/R(X +1)) 1.288-273.15 ⇒RX:JMP (X+5→X)> BH 21: 80→A;4→C;0→X; LTR R97+.5,R25,2 11;PLT "TEMP-DRY 22:  $R(A+X) \rightarrow R(A+X-1)$ JMP (X+1+X)>CF23: A-5+A:C+5+CH 24: IF A≠0;0→X;GTO 2 21 25: 48 (5+3+0 #R (C-A)+ Ric-A+4);JMP (A+ 1 + 0 1 > 0 F 26: 4+%;C+R0;R99+R1; R98→R2:ENT "DS", R3:REC "DA",R80: DSP "LOD OPT"; STP F 27: IF 0>R(X+2);GTO 295 28: PLT R(X+2), RXH 29: IF RODX:X+4+X; GTO 27H 30: 4⇒X;LTR R18+.5,R 16.211; PLT "WIND SPEED"; STP F 1112 PLI R(M+3), RM; JMP (X+4+X)>R0H 33: 0+Y%LTR R47+.5,R 44,211; PLT "POTE NTIAL" \$ LTR R47+1 .5,R44-75,211; PLT "TEMP"H 33: PLT R7,Y;PLT R7, 25+Y;PEN ;JMP (Y +50+Y)>2050F 34: R7+.5→812000+8H 35: LTR A, B, 211; PLT "ADIABATIC POTEN TIAL"; LTR A,8-75,211; PLT "TEMP. GRAD."H

36: STP ;5→X;0→R88;3 60÷R89H 37: IF RX>360;GTO 40 38: IF RX>R88;RX+R88 39: IF R89>RX;RX+R89 40: IF R0>X;X+4→X; GTO 37H 41: FXD 0H 42: 2INT (R88/40)+2+ AH 43: 2INT (R89/40) - BH 44: A-B+YH 45: 36-Y+0;0+XH 46: LTR 6+2X-.8,4370 ,211:PLT 20B+40X \$JMP (X+1+X1)Y-2 47: 4 + X |-48: IF RtX+1+>360; 6T0 50H 49: PLT 36-(20A-R)X+ 11)/20, RXF 50: IF RO>X;X+4+X; GTO 48F 51: LTR 32.5-120A-R2 51/20,R24,211; PLT "WIND DIR"H 52: FMD 1;LTR 2,2225 ,211;PLT "T-DRY-":PLT R1:PLT " C |--53: LTR 2,2150,211; PLT "P-";PLT R2; PLT 54: ENT "MONTH", R98. "DAY" #R91, "YFAR" ,P92,"TIME",R93F 55:

56: LTR 15,2310,211; PLT R90;PLT "/"; PLT R91;PLT "/"; PLT R92; PLT ", ; PLT R93; PLT "Z" 57: LTR 2,2075,211; PLT "DEN-";PLT R SIPLT " GMZMS"H 58: SCL ~75,1000,-20 0,3100;FXD 0; TBL 2;DSP "C-GRA PH";STP + 59: R1+273.15→R1:4→X ;0→B;CFG 1;1→Z; PLT 0,0;CFG 2F 60: P1R3+6.145500432 E - 12 + R(X + 4) - B) + 3.5163→A⊦ 61: (9.8/P1)(R()+7)-BR82-R71/1R(X+4) -B-R4+1E-8) →CH 52: IF 1E-6 0; ∩A→Y; SFG 2:070 66F 63: IF 1-At 3/-1; ACS (1 AC/2)//C→ YIGTO GAISTE H 64: (PIX+ | PIX+3)) (R(X+4)- PX++R82F 651 B+10→8: SFG 1; GTO 66F 66: R18+41-B+R861 PLT Y.FOSTIF FLG 1=1:PLT 900; R86; YəF85; 901əYF 67: IF Y≤90057+R(64+ Z) ; X+4+; ; 1+Z+Z; GT0 600 68: LTR 9:0:006-20:2 11; PL1 " "; PLT R86; R85 > A (64+Z); Z≯R64F 69: LTR 660, RS6-100, 211% PLT "STABLE ATMOSFHERE" | STP

1

FXD 0F

## TABLE 2 (Concluded)

```
70:
R(X+4)-R(X-4)+R7
61(R(X+6)-R(X-2)
))/R76)(R(X+4)-B
-R(X-4))+R(X-2)+
Rash
7 1 8
(R_1X+5) R(X-3)
ZR76);RX-8-RIX-4
J:+R:(7-::;-270+R8
21-
72:
LTR 9:2800:2111
FLT "hE".DATA-"!
PL1 R90 PLT "/ ;
TE" POLIPET 12 4
PLI 992 PIT " '4
FUT ROSSPIT TO A
. ] #
FRU 137 FR 982300
821746 T 157ME 1
F RIEW CIPLE FBS
i ...
1 4 1
IF FLG REIGHTR ...
0,100 211101 F P
DIABATIL ATMOSPH
后原用"补
11...
1.1111 -
F105
```

### TABLE 3. LISTING OF DIFFUSION ROUTINE

ENT "OT", R79, UT ", R81: 2.18165562 10E10R1/R2+R78;0 →R80;0→X;FXD 0H 1 5 ĒNT "0AR",R83,"0 ATK",R84,"HL",R8 7;.26512R86→R89H SCL -3000,20500, -100,970;-20000+ R102+R103;1+R88H 3: LOG (R84/R83)/ LOG (R87/2)→CH 4: (R85/600)1.2πR83 (R87\*(C+1)-2\*(C+ 1) 1/1 (C+1) + (R87-2)+2↑C\*180:→R101 E .. CLL FA X<sub>2</sub>R97<sub>3</sub>R77 ,R1,BH 6: R78B/(2mR77\*R2r( A/2):⇒Y;R78B/(R7 7\*# (A/2)2## (2#)R 1) →AF 7: PLT X-500, R102; PLT X-100A; PEN ; 100A⇒R102;PLT X-500,R103;PEN ; PLT .:, YIPEH ; Y+R 103F 8: ĪF A RSW;A→R80;X \*R98E 9: IF 20000) X 9 X + 500 ⇒X;GYO 5⊦ 10: LTR 300,860,211; PLT R90;PLT "/"; PLT R91;PLT '/"; PLT R92;PLT " "; PLT R93;PLT "Z"F 11: FXB 1:LTR 300,92 0,211;PLT "HEIGH T OF LAYER-'; PLI ROTAPLY " M" 12: PRT R79-R51

13: LTR 17000,860,21 1; PLT "CONC. MAX-";PLT R80;4→X; CFG 0; CFG 1; CFG 2;DSP "ISO-GRAPH "iSTP F 14: FXD 1; IF R(X+2)> 0;GTO 17H 15: R(X+4)-R(X-4)+R76i((R(X+6)-R(X-2))):/R76)(RX-R(X+4 ))+R(X-2)+R(X+2) 16: ((R(X+5)-R(X-3)) $\angle R76)(RX-R(X-4))$ +R(X-3)+R(X+1)F17: IF 60>X;X+4→X; GT0 14F 18: SCL -20000,18000 ,-24000,6050; TBL 1F 19: 4+X;R64+R97;0+A+ B+Y+R64F 20: R(X+2)(R(65+B)-R(64+B))→C;R(X+1) -270+ZH 21: CCOS Z+Y+Y;CSIN Z÷A⇒A;PLT A,YE 22: lF R97-1>B;B+1→B ;X+4→X;GTO 20H 23: IF FLG 2≠1;A→R76 >R94;Y>R77>R95; SFG 2H 24: C+60R99~C;R82+Z; IF 4000>C;GTO 21 25: LTR R76-50, R77-9 0,211;PLT "X ":PLT R85+R96H 263 IF FLG 1≠1;A/R99 9 FM | ZHP961A4R761 Y>R77; SFG 1; GTO 25F

27:
R94+R985:N R82→A
;R95+R98cc5 R82→
B;LTR A,B,211;
PLT "O"H
28:
LTR 500,-20000,2
11;PLT R85+R98/R
99;PLT "SEC"H
29:
LTR 500,-21000,2
11;PLT F(A↑2+B↑2
);PLT "M ";
PLT ATN (B/A);
PLT "DEG"H
30:
FXD 0;LTR -19500
,-21000,211;PLT
R90;PLT "/";PLT
R90;PLT "/";PLT
R92:PLT "":PLT
R93;PLT ZH

311 ENT "C", R88 + 2; -4 0000+R17+R18;0+X 32: CLL FA WAYE 33: 1000+R21H 34: CLL: FA R21 - R20+ 35: X-Y(R21-X)/(R20-Y)→R22F 36: IF 20>ABS R20; ABS Y⇒Y∜GTO 39E 37: CLL FA R22, R23H 38: X+R21;Y+R20;R22+ X;R23⇒Y;GTO 35⊬ 39: ΓY⇒Y;R95+XCOS R8 2-YSIN R82→R50;R 95+XCOS R82+Y SIN R82+R51F 40: R94+YC08 R82+X SIN R824R52;R94-YCOS R82+XSIN R8 2+R53F 41: PLT R9, R17; PLT P 52,R50;PEN ;R52, R9;R50-R17F 42: PLT R8, P18; PLT P 53,R51;PEN ;R53+ R8(R51→R13()+ ABS (R17 R)8+10) /3∌%E 43: IF ABS PA 18000; GTO 30H 44# IF R17) bod0 # GTO 311 45% IF -22000, R17; GTO 31H 46: FXD 1) IF FLG Ø=1 \$LTR R52,R50,211 ;PLT P88:CFG 0+ CLL FA L. YE 48: GTO 350 49: END + RI 08

## TABLE 3 (Concluded)

```
ill :
0.8:2(P1R101+1 P)
1 :
IF 12288718(6) N.
78>325:GTO 45
. .
SERR CHIZZROZ-RE
KITE HITTER IN TO
22R87. P861 P2/4) +
RyBE
3.5
Z+1-24010 1F
深起图图 (1) 斯马格莱德 (1902)。
Dadable Harrist
111
Compatibility of the second
· 医克里克氏 化二甲基甲基二丁
distribution of the
Single Alberta Berlin
LUG TRADE BOOK
MILLESON CLAYINGS
241-25
The Book of the Artist
1 1
d - 3, 1-
The second
· 益聚合物 (114) (4046)。
推开的1200年中国1916年2月中
工程 10.1 人名英格兰人名
21
:2:
KE) H
13:
ENU H
F 111
```

## FA SUBROUTINE

# TABLE 4. LOCATION OF EQUATIONS IN THE NASA/MSFC SURFACE-DIFFUSION PROGRAM (MET ROUTINE)

## CLOUD RISE

Instantaneous Adiabatic Atmosphere Line 60

$$t^2 = z^{3.5163}$$
 PT × Const.

Const. = 
$$\frac{2 \pi C_p V^3}{3 g Q}$$

Instantaneous Stable

Line 60, 63

$$t = COS^{-1} \left( 1 - \frac{z^{2.5163} PTS \times Const.}{2} \right) / \sqrt{s}$$

Const. = 
$$\frac{2 \pi C_p V^3}{3 g Q}$$

Potential temperature gradient

Line 61

$$\mathbf{s} = \frac{\mathbf{g}}{\mathbf{T}} \frac{\Delta \Phi}{\Delta \mathbf{z}}$$

# TABLE 5. LOCATION OF EQUATIONS IN THE NASA/MSFC SURFACE-DIFFUSION PROGRAM (DIFFUSION ROUTINE)

Line 0

$$Q_{K} = F \left(\frac{10^{3} \text{ mg}}{\text{g}}\right) \left(\frac{22.4}{\text{M}}\right) \left(\frac{T\{z_{R}\}}{273.16}\right) \left(\frac{1013.2}{P\{z_{R}\}}\right)$$

Line 1

$$\sigma_{xo}$$
,  $\sigma_{yo}$ ,  $\sigma_{zo} = \frac{\gamma H}{4.3}$ 

Line 3

$$m = \log \left( \frac{\sigma_{ATK}^{\prime}}{\sigma_{AR}^{\prime}} \right) / \log \left( \frac{z_{T}}{z_{R}} \right)$$

Line 4

$$\sigma_{AR}^{\prime} = \sigma_{AR} \left(\frac{\tau_{K}}{\tau_{oK}}\right)^{1/5} \left(\frac{\pi}{180}\right)$$

$$\sigma_{A}^{\prime} = \frac{\sigma_{AR}^{\prime} (z_{T})^{m+1} - (z_{R})^{m+1}}{(m+1) (z_{T} - z_{R}) (z_{R})^{m}}$$

Line 6

$$D = \frac{Q_{K}}{2\pi \sigma_{y} \sigma_{z} \bar{u}_{K}} \left(2 \exp \left[\frac{-(H)^{2}}{2\sigma_{z}^{2}}\right] + \sum_{i=1}^{\infty} \left\{2 \exp \left[\frac{-(2i(z_{T} - H))^{2}}{2\sigma_{z}^{2}}\right]\right\}$$

# TABLE 5. (Concluded)

+ 
$$2 \exp \left[ \frac{-(2i(z_T + H))^2}{2\sigma_z^2} \right]$$
  $\chi = \frac{D \bar{u}_K}{\sqrt{2\pi} \sigma_x}$ 

# TABLE 6. LOCATION OF EQUATIONS IN THE NASA/MSFC SURFACE-DIFFUSION PROGRAM (FA SUBROUTINE)

Line 0

$$\sigma_{z} = \sigma_{E}^{\dagger} x + x_{z}$$

Line 1, 2, 3

$$\sum_{i=1}^{\infty} 2 \exp \left\{ \frac{-[2i(z_{T} - H)]^{2}}{2\sigma_{z}^{2}} \right\}$$

$$+ 2 \exp \left\{ \frac{- [2i (z_{T} + H)]^{2}}{2\sigma_{z}^{2}} \right\}$$

Line 4

$$2 \exp \left[ \frac{-(H)^2}{2\sigma_Z^2} \right]$$

Line 5

$$\sigma_{y} = \left[ \left( \sigma_{A}^{\prime} x + x_{y} \right)^{2} + \left( \frac{\Delta \theta^{\prime} x^{2}}{4.3} \right)^{2} \right]$$

Line 6

$$p = \log \left( \frac{\bar{u}_{T}}{\bar{u}_{R}} \right) / \log \left( \frac{z_{T}}{z_{R}} \right)$$

Line 7

$$\tilde{u}K = \frac{\tilde{u}_{R} \left[ (z_{T})^{1+p} - (z_{R})^{1+p} \right]}{(z_{T} - z_{R}) (z_{R})^{p} (1+p)}$$

Line 8, 9

$$L\{x\} = \frac{0.28 (\Delta \bar{u}_{K}) (x)}{\bar{u}_{K}}$$

Line 10

$$\sigma_{X} = \left[ \left( \frac{L\{X\}}{4.3} \right)^{2} + \sigma_{XO}^{2} \right]^{\frac{1}{2}}$$

Line 11

Concentration Solved For y Position.

TABLE 7. AVERAGE VALUES FOR SURFACE LAYER DIFFUSION ROUTINE

$Q_{K}$	6.19.10 <sup>9</sup>			
$\sigma_{yo}$ , $\sigma_{xo}$ , $\sigma_{zo}$	194.7			
m	-0.0979			
$\sigma_{\mathbf{A}}^{\dagger}$	-0.07203			
Downrange distance	0	500	1000	1500
$\sigma_{z}^{2}/z$	7.58.10 <sup>4</sup>	1.056.105	1.430.103	1.8.10 <sup>5</sup>
Σ	0	2.10-5	4.10-4	2.61.10-3
vert. term	.00163	.01215	.04329	.10263
$\sigma_{ m y}$	194.7	261.03	362.61	477.4
P	0	0	0	0
$\overline{\mathtt{U}}$	6.18	6.18	6.18	6.18
L(X)	0	0	0	0
$\sigma_{ m X}$	194.7	194.7	194.7	194.7

### TABLE 8. LABELING ROUTINES FOR GRAPHS

SCL -4,36.5,-400 ,2400;AXE 0,0,1, 100;AXE 0,-250,1 ,0;AXE 36.5,2400 ,1,100H i : LTR 1,2300,211; PLT "SURFACE CON DITIONS"; FXD 0H 2: 0⇒A;0⇒BH 3: LTR -2, A-25, 211; PLT A; JMP (A+200 +A) >2000F 4: LTP -3:600,222; PLT "ALTITUDE-ME TERS" H 5: LTR B-.2,-60,211 ;PLT B; JMP (B+2+ B)>36H File LTR 1,-140,221; PLT "VELOCITY M/ SEC---TEMPERATUR E-DEG C"H 7: LTR 10,-400.221; PLT "WIND DIRECT ION-DEGREE"+ 8: LTR 28,2300,211; PLT "PAWINSUNDE DATA"E 9: LTR 28,2150,211; PLT "CAPE MENNED Y, FLA."H 10: LTR 14.2300.211% PLT "S+E AERO-YA "⊢ 11: LTR 14,2225,211; PLT "NASA / MSEC " |-12: LTR 14,2150,211; PLT "(00T. 73,P. А.Н.)"⊦ 13: END H R340

из SCL -75,1000,-20 0,3100;AXE 0,0,3 0,100;FXD 0;0→XH 1: LTR 9,2900,221; PLT "TEMPORAL AS CENT OF A TITAN III EXHAUST CLOU D' F 2: LTR 270,2800,211 ;PLT "CAPE KENNE DY: FLA."H 3: LTR 9,2600,211; PLT "S+E-AERO-YA NASA / MSFC"H 4: LTR 9,2500,211; PLT "(OCT.73, P. A.H.) "F 5: LTR -60,950,222; PLT "ALTITUDE-ME TERS"H 6: LTR -50,X-20,211 FPLT XJJMP (X+20 U→X)>3000H 7: LTR 225,-160,221 ;PLT "TIME AFTER LAUNCH-SECONDS" ;0→XH 8: LTR N-10,-80,211 FPLT XFJMP (X+60 ⇒X)>900F 약 : END H R354

SCL -3000,20500, -100,970;AXE 0,0 ,1000,50;AXE -15 00,0,0,50H 1: LTR ~2700,300,22 2; PLT "CONCENTRA TION-PPM"; FXD 2; 0+XE 2: LTR -2500,X-3,21 I;PLT X/100;JMP (X+50+X)>950F 3: FXD 0;LTR 7000:-70,221;PLT "DIST ANCE-KM";0→XH 4: LTR 1000X-100,-3 0,211;PLT X;JMP (X+1+X)>20F5: LTR 300,950,221; PLT "CENTERLINE CONCENTRATION AN D DOSAGE"H 6: LTR 300,830,211; PLT "NASA / MFSC S+E-AERO-YA"H LTR 300,800,211; PLT "(OCT. 73, P .A.H.)"h 8: LTR -1000,320,22 2 PLT "DOSAGE-PP M-SEC";FXD 0;0→X 9: LTR -800,X-3,211;PLT X;JMP (X+50 →X)>950H 10: LTR 17000,920,21 I PLT "CONCENTRA TION-";PLT 20200 ,920;PLT 20500,9 201 11: LTR 17000,890,21 1;PLT "DOSAGE-. 12: LTR 300,890,211; PLT "RAWINSONDE DATA"E 13:

END H

SCL -20000,18000 ,-24000,6050; AXE 0,0,1000,100 0;0→X;FXD 0⊢ 1: PLT 6000COS X,60 00SIN-X; JMP (X+1 ⇒X)>360H 0+A;0+BH 3: PLT (5900+8)COS A, (5900+B)SIN A; JMP (B+200∍B)>20 ØΕ 4: PEN ; IF A≠360; A+ 10→A;0→B;GTO 3⊢ 5: LTR -19500,-1900 0,221;PLT "TITAN III HCL"F LTR -19500,-2000 0.211;PLT "RAWIN SONDE DATA"E 7: LIR -19500,-2200 0,211;PLT "NASA / MESC S+E-RER 0-7A"E 8: LTR -19500;-23и0 0:211;PLT ":ОСТ. 73. P.A.H.!"F 9 : LTR 500.-19000.2 11:PLT "CONC. MA X. CONDITIONS"F 10: END H P354

\_ |

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